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1	Hydrological classification of natural flow regimes to support
2	environmental flow assessments in intensively regulated
3	Mediterranean rivers, Segura River Basin (Spain)
4	
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15	Abstract
16	Hydrological classification constitutes the first step of a new holistic framework for
17	developing regional environmental flow criteria: the "Ecological Limits of Hydrologic
18	Alteration (ELOHA)". The aim of this study was to develop a classification for 390
19	stream sections of the Segura River Basin based on 73 hydrological indices that
20	characterize their natural flow regimes. The hydrological indices were calculated with
21	25 years of natural monthly flows (1980/81 - 2005/06) derived from a rainfall-runoff
22	model developed by the Spanish Ministry of Environment and Public Works. These
23	indices included, at a monthly or annual basis, measures of duration of droughts and
24	central tendency and dispersion of flow magnitude (average, low and high flow

1 conditions). Principal Component Analysis (PCA) indicated high redundancy among 2 most hydrological indices, as well as two gradients: flow magnitude for mainstream 3 rivers and temporal variability for tributary streams. A classification with eight flow-4 regime classes was chosen as the most easily interpretable in the Segura River Basin, 5 which was supported by ANOSIM analyses. These classes can be simplified in 4 6 broader groups, with different seasonal discharge pattern: large rivers, perennial stable 7 streams, perennial seasonal streams and intermittent and ephemeral streams. They 8 showed a high degree of spatial cohesion, following a gradient associated with climatic 9 aridity from NW to SE, and were well defined in terms of the fundamental variables in 10 Mediterranean streams: magnitude and temporal variability of flows. Therefore, this 11 classification is a fundamental tool to support water management and planning in the 12 Segura River Basin. Future research will allow us to study the flow alteration-ecological 13 response relationship for each river type, and set the basis to design scientifically 14 credible environmental flows following the ELOHA framework.

15

Keywords: Ecological Limits of Hydrologic Alteration (ELOHA) · Environmental
 flows · Regulated Mediterranean rivers · Modelled monthly flows · Temporal
 variability · Intermittent streams · Drought

19

#### 20 Introduction

21

Flow regime has become a fundamental part of running water ecosystems ecological studies and management (Arthington and Pusey 2003; Bunn and Arthington 2002; Richter and others 2006). Since the publication of the "natural flow regime paradigm" (Poff and others 1997), ecologists have recognized intra and inter-annual flow variability as a primary driver of the structure and function of riverine ecosystems and

many of the adaptations of its biota (Arthington and others 2006; Lytle and Poff 2004;
Naiman and others 2008). Many authors have emphasized the need to characterise the
similarity among flow regimes to provide typologies that can support *a priori*predictions (e.g. ecological and evolutionary convergence under geographically disjoint
regimes) and the development of general principles for flow regime management, such
as the assessment of environmental flows (Arthington and others 2006; Poff and others
2006).

8

9 "Environmental flows" is now a widely accepted term that covers the "quantity, timing, 10 duration, frequency and quality of water flows required to sustain freshwater and 11 estuarine ecosystems and the human livelihood and well-being that depend on these 12 ecosystems" (Brisbane Declaration, 2007). Implementing environmental flows will be a 13 key measure for protecting and restoring river ecosystems (Arthington and others 1991; 14 Arthington and others 2010; Poff and others 1997; Richter and others 1996; Richter and 15 others 1997; Sparks 1992; Stanford and others 1996). More than 200 methodologies 16 have been described to define environmental flows since the decade of 1970 (Tharme 17 2003). Hydrological methods are based on the study of long hydrological series. The 18 simplest ones only define rules to set a minimum flow for the river (Tennant 1976). 19 However, there are more complex approaches such as the RVA (Range of Variability 20 Approach) method (Richter and others 1997). This method characterizes flow records 21 using 32 different hydrological parameters, the "Indicators of Hydrologic Alteration", 22 and establishes a range of variation (for example, the mean  $\pm$  the standard deviation) as 23 the objective for each one. The authors point out that these objectives must be 24 completed for every river with field research, suggesting that hydrological records 25 cannot be the only source of information in the definition of environmental flows.

Habitat simulation methods determine the flow requirements of some "target species", 1 2 usually fishes (Bovee 1982), but they have been applied to macroinvertebrates (King 3 and Tharme 1994) or even to achieve objectives related to the morphology of the river 4 (Milhous 1998). Finally, holistic methodologies broaden the definition of 5 "environmental flow" considering the fluvial ecosystem as a whole instead of focusing 6 only in the requirements of a few species (Arthington and Pusey 1993; King and 7 Tharme 1994; Poff and others 1997; Richter and others 1996; Sparks 1992; Sparks 8 1995). The relationship between flow alteration and ecological characteristics for 9 different river types constitute the key element of a new holistic framework for 10 developing scientifically-credible regional environmental flows criteria: the "Ecological 11 Limits of Hydrologic Alteration" (ELOHA) (Arthington and others 2006; Poff and 12 others 2010). A principle for setting environmental flows is that this should be carried 13 out at a regional level, because they are related to river types that will have differing 14 natural or "reference" conditions (Poff and others 2010). Therefore, there is a need to 15 develop river classifications to identify the natural flow regime for each stream, to 16 develop the flow-ecology relationship and to assist the assessment of environmental 17 flows.

18

Several hydrological classifications have been made for large river basins (Hannah and others 2000; Harris and others 2000), states (Apse and others 2008; Cade 2008; Kennen and others 2007; Kennen and others 2009) or even entire countries, such as USA (Mcnamay and others 2011; Olden and Poff 2003; Poff 1996), New Zealand (Snelder and Biggs 2002), Germany (Pottgiesser and Sommerhäuser 2004), France (Snelder and others 2009), Australia (Kennard and others 2010) and Chile (Peredo-Parada and others In Press) using different methods. Two basic approaches have been used to achieve this goal: (1) *a priori* classifications using climatic and other environmental variables that
 influence hydrology and (2) *a posteriori* classifications based on hydrological statistics.

3

4 In Spain, according to the water legislation (revised text of the Water Law, 2001), 5 environmental flows should be included in Basin Management Plans to fulfil the EU 6 Water Framework Directive (WFD, 2000). However, no national hydrological 7 classification has been published. Ecoregions and ecotypes classifications based on non-8 altered geographical, morphological, climatic and geological variables have previously 9 been attempted following the WFD system B (Anex II) at national (CEDEX 2004) and 10 Mediterranean scale (Bonada and others 2002; Moreno and others 2006; Munne and 11 Prat 2004; Sanchez-Montoya and others 2007), respectively. But these classifications 12 did not include hydrological variables or described only one or two flow-regime 13 components (e.g. mean annual discharge). Nevertheless, hydrological classifications 14 based on hydrological indices have been developed for the Tajo and Ebro basins 15 (Alcazar and Palau 2010; Baeza and Garcia de Jalon 2005; Bejarano and others 2010).

16

The present study addresses a hydrological classification for stream and river segments in the Segura River Basin, an intensively regulated Mediterranean basin in the Southeastern Spain, based on the similarity in their natural flow regimes, characterised using hydrological indices. Specific objectives were to determine the hydrological variables that best discriminate and characterize the different flow types and to identify the spatial distribution of the resulting river classes.

23

24 Methods

2

3 The Segura River Basin, as management unit (including coastal watercourses), 4 represents one of the most arid zones of the Mediterranean area, presenting great 5 heterogeneity in its flow regimes. It is located in the SE of Spain (Fig. 1), one of the 6 most arid zones of the Mediterranean area in this country. Despite its small size (18870 7 km<sup>2</sup>), there is a strong climatic and altitudinal gradient from NW to SE. The climate 8 ranges from wet (>1000 mm mean annual precipitation) and cold in the mountains 9 (>1000 masl) of the NW to semiarid (< 350 mm mean annual precipitation) in the SE 10 lowlands (200 mm precipitation near the coast). Mean annual temperatures range 11 between 10 and 18 °C (CHS 2007). The lithology of the plains is characterised by the 12 dominance of limestone as well as Miocene and Triasic marls, with some volcanic 13 areas, whereas calcites and dolomites dominate the mountain headwaters. The landscape 14 ranges from Mediterranean conifer forests in the mountains to arid and semi-arid 15 shrublands in the south-east lowlands. This longitudinal gradient in altitude and climate 16 is coupled with a human density gradient. The river network has low populated forested 17 headwaters, populated agricultural midlands with intense flow regulation and densely 18 populated cities in the lowlands (Mellado 2005). Agricultural (52.1%), forest and semi 19 natural (45.2%) and artificial (2.1%) land uses predominate in the Segura Basin 20 (estimated from Corine Land Cover 2000).

21

As for other Mediterranean regions, the basin is characterised by scarce and unevenly distributed water resources and high hydrologic variability (low rainfall irregularly distributed in time and space). Large storm events often produce flooding during spring and autumn (CHS 2007). High temperatures and low rainfall during the summer season

1 lead to a natural water scarcity, generating drought events and in some cases the complete cessation of flow. The largest volume of surface water is provided by the 2 3 tributaries in the upper sector of the basin. The Mundo River, the major tributary, 4 provides most of water resources. The regulation capacity by dams (24 dams higher than 10 m.) in the Segura Basin is approximately 770 hm<sup>3</sup>, equivalent to over 90% of its 5 6 natural input (CHS 2007). There is also significant regulatory volume (approximately 7 325 hm<sup>3</sup>) of inter-basin transfers from the Tagus River. Mean groundwater abstraction 8 is 478 hm<sup>3</sup>/year, over 80% of the natural recharge. Water for irrigation represents the 9 main water withdrawal (90% of resources). These human activities in the rivers and 10 their catchments profoundly alter the natural flow regimes, producing a significant 11 reduction in the magnitude of flows and a reversal in their seasonal pattern. River 12 reaches below dams present maximums in summer and minimums in winter, with 13 droughts becoming more frequent and long-lasting (Vidal-Abarca and others 2002; 14 Belmar and others, 2010).

15

16 Drainage network

17

A drainage network was derived from a 25 m. digital elevation model (DEM), developed by the National Geographic Institute of Spain (IGN, 2005), and fragments extracted from layers available in the website of the Ministry of Environment, in order to achieve higher precision. The ArcGIS software (v 9.2) with the ArcHydro extension (v 1.2) (ESRI, Redlands, California, U.S.A., 2006) were the tools used. The network comprises sections that link each network junction (node). Each node, at the end of each section, is associated with its corresponding watershed (derived from the DEM). The minimum watershed area to define a section was 10 km<sup>2</sup>. The hydrological network
comprises 390 nodes and sections (Fig. 1).

3

4 Baseline or reference flow conditions

5

Within the Segura Basin, there is limited hydrological information from gauging
stations representing unaltered regimes. Gauged sites are scarce and located principally
in the mainstream; impacted by dam and reservoir operations, water withdrawals and
diversions.

10

11 To build a database of flow time-series that represents the baseline or reference 12 conditions we used the SIMPA model (the Spanish acronym meaning "Integrated 13 System for Rainfall-Runoff Modelling"), developed by the Centre for Hydrographic 14 Studies (CEDEX, Ministry of Environment and Public Works, Spain). This model is an 15 implementation of a classic soil moisture balance model (Temez 1977) where soil and 16 aquifer storages are considered, as well as a collation of transfer laws (Estrela and 17 Quintas 1996a; Estrela and Quintas 1996b; Ruiz 1998). Some publications illustrate 18 SIMPA's progress (Alvarez and others 2005; Barranco and Alvarez-Rodriguez 2009; 19 Potenciano and Villaverde 2009). It takes monthly precipitation from 1 km. grid maps 20 created by the Spanish Ministry of Environment by means of an interpolation procedure 21 (the inverse to the square distance) with data from the more than 5000 weather stations 22 of the Spanish network. For this interpolation, double regression and "white noise" 23 procedures were used to complete incomplete series without altering the natural 24 variance of data, as well as specific procedures for the highest elevation areas (Estrela 25 and others 1999). Calibrated by regionalization of different variables (maximum moisture capacity, as a function of land use; maximum infiltration, as a function of
lithology; etc.), the model has been validated by means of comparison with reference
and restored records in more than 100 control points (Estrela and others 1999). Besides,
it has been used in Spain for water resources assessment, in the White Paper Book of
Waters (Ministry of Environment 2000) and the National Water Master Plan (Ministry
of Environment 2002), and for a hydrological classification of the streams and rivers in
the Ebro Basin (Bejarano and others 2010).

8

9 We generated monthly data that represented natural flow conditions for the period
10 1980/81-2005/06 in each node of the hydrological network to calculate a set of
11 hydrological indices.

12

13 Classification of river flow regimes

14

15 73 hydrological indices describing either monthly or annual characteristics (see 16 Appendix) were calculated. These indices, based on the "Indicators of Hydrologic 17 Alteration" (Mathews and Richter 2007), represent a wide range of ecologically-18 relevant flow statistics (Mathews and Richter 2007; Monk and others 2006; Monk and 19 others 2007; Olden and Poff 2003; Richter and others 1996) and include measures of the 20 duration of droughts as well as the central tendency and dispersion of flow magnitude 21 (average, low and high flow conditions), two of the major components of the flow 22 regime in Mediterranean rivers. However, other significant components related to the 23 frequency, duration and rate of change of high flood events were not estimated because 24 of the lack of daily flow data.

1 Hydrological indices have considerable multicollinearity (Olden and Poff 2003). We 2 reduced our set to a smaller set of non-redundant indices using the procedure outlined in 3 Olden and Poff (2003). A principal components analysis (PCA) was used to examine 4 dominant patterns of intercorrelation among the hydrological indices and to identify 5 subsets of indices that describe the major sources of variation while minimizing 6 redundancy (i.e. multicollinearity). This PCA was conducted, using PC-ORD (v 4.41) 7 (McCune and Grace 2002), with the correlation matrix rather than the covariance matrix 8 to ensure that all indices contributed equally to the PCA and that these contributions 9 were scale-independent (Legendre and Legendre 1998). We selected the simplest and 10 most easily interpretable indices to characterize flow regimes, based on criteria of high 11 correlation with the three first PCA axes.

12

13 Scores for the first three axes were weighted by the proportion of the variance explained 14 by each PCA axis and used as new synthetic hydrological variables for a cluster 15 analysis. A flexible- $\beta$  clustering technique (Legendre and Legendre 1998; McCune and 16 Grace 2002) was used to group streams according to their similarity in flow regime, 17 measured using Euclidean distances. This technique allows the user to select the number 18 of clusters desired and choose the most interpretable classification. Besides, as an 19 internal validation, Analyses of Similarities (ANOSIM) (Clarke 1993) were run on the 20 Euclidean distances to test the effect of the number of classes on the degree of 21 separation among them. Each test in ANOSIM produces an R-statistic, which contrasts 22 the similarities of nodes within a class with the similarities of nodes among classes 23 (when the R value is close to one, similarities between nodes within a class are higher 24 than those between nodes from different classes, and values close to zero indicate no

differences among classes). These analyses were conducted in PRIMER (v 6) (Clarke
and Gorley 2006).

3

In order to visually appreciate the differences between hydrological classes we represented annual hydrographs showing the standardized monthly flows of the streams and rivers included in each class, as well as whisker box plots showing environmental variables: average precipitation in the drainage area, drainage area, Strahler order (Strahler 1957), average altitude as well as slope of the drainage area and percentage of karstic surface. The latter was derived from the Spain's Map of Karst (1:1.000.000) developed by the Mining Geologic Institute of Spain (IGME, 1986).

11

13

14 Redundancy among hydrological indices

15

Most of variation (73.35%) in the hydrological variables was explained by the first two axis of the PCA. Figure 2 presents the two-dimensional ordination illustrating the major patterns of inter-correlation among the 73 hydrological indices for the combined set of 390 stream and river sections; the symbols by stream classes correspond to the clusters, shown in Figure 3. The majority of indices were highly correlated (either positively or negatively). The percentage of months with zero flow ( $D_L$ ) was the only one with a high significant correlation with all the other indices.

23

Three groups of hydrological indices were differentiated. A first group, in the first quadrant of Figure 2, included indices related to the intensity of droughts  $(D_L)$  and 1 floods, such as indices of annual maximums (AMAX/Q50,  $I_H$ ). This group also included 2 indices of dispersion describing the variability of the flow regime, such as the 3 coefficient of variation in mean annual flows ( $CV_{INTER}$ ), the coefficient of variation in 4 mean monthly flows ( $CV_{INTRA}$ ), the coefficients of variation in monthly flows ( $CV_A$  1-12), the coefficient of variation in maximum monthly flows ( $CV_H$ ) and other variability 6 indices based on percentiles (Q5/Q50, Q10/Q50).

7

8 In the third quadrant there was a second group of indices. This group contains indices 9 that characterize the magnitude of low flows, such as the mean minimum monthly flows 10  $(M_L 1-12)$ , the average of minimum monthly flows  $(M_L 13)$ , the annual minimum 11 discharge divided by the median (AMIN/Q50); and the magnitude of average flows, 12 such as the mean and median annual runoff  $(M_A 16$  and MEDDIS/A).

13

A third group of correlated variables (second quadrant) included measures of central tendency in flow magnitude and high flows, such as the mean and median annual discharge (MADIS, Q50), mean monthly flows (M<sub>A</sub> 1-12), mean maximum monthly flows (M<sub>H</sub> 1-12), the average of maximum monthly flows (M<sub>H</sub>13) and some measures of variability (STDEV, Q1, RANGE).

19

From the non-correlated indices in the two first quadrants, the mean annual discharge (MADIS), the percentage of months with zero flow ( $D_L$ ) and the coefficient of variation in mean annual flows ( $CV_{INTER}$ ) represent the major gradients of variation in the Mediterranean flow regimes. The two first indices were highly correlated (negatively and positively, respectively) with the first axis, while  $CV_{INTER}$  was correlated with both PCA axes. Thus, stream and river sections were interpreted in the two-dimensional space (Fig. 2) following two gradients: (1) a flow magnitude gradient, crossing the second quadrant, that ordered the mainstream sections of the rivers Segura and Mundo from larger (upper left corner) to smaller discharge; and (2) a temporal variability gradient, crossing the first and third quadrant, that ordered the tributaries from ephemeral and intermittent streams (upper right corner in Fig. 2) to permanent and more regular ones.

7

8 Hydrological classes

9

10 With the  $\beta$ -flexible clustering based on weighted PCA scores, a classification with eight 11 hydrological classes (Fig. 3) was chosen as the most easily interpretable solution for the 12 Segura River Basin. Besides, the ANOSIM analyses defined the 8 classes solution as the 13 most convenient. It produced the greatest increase in the R-value and, despite that the 9 14 classes solution produced the biggest R-value, the increase is negligible (Fig. 4). The 15 magnitude of annual flows (MADIS), the duration of droughts (D<sub>L</sub>) and the interannual 16 variation of flows (CV<sub>INTER</sub>) were discriminators of these 8 flow-regime classes (Fig. 5). 17 The first division of the cluster distinguished between perennial mainstream rivers (Classes 1–2), with an average annual flow larger than  $2 \text{ m}^3/\text{s}$ , and tributaries (Classes 3-18 19 8), with smaller mean discharges. Tributaries include sites ranging from perennial 20 streams, which never (Classes 3-4) or eventually (Classes 5-6) cease flowing, to 21 intermittent and ephemeral streams (Classes 7 and 8), which stop flowing a 20% and a 22 50% of time respectively.

23

Therefore, the eight classes (Fig. 3, bottom dotted line) can be grouped into four broader groups (Fig.3, upper dotted line): *large rivers* (Classes 1 and 2), *perennial stable* 

streams (Classes 3 and 4), perennial seasonal streams (Classes 5 and 6) and intermittent
 and ephemeral streams (Classes 7 and 8). Distinctions within each couple were evident
 in terms of differences in annual hydrographs (Fig. 6) and environmental characteristics
 of the watersheds (Fig. 7).

5

6 Classes 1 and 2, perennial large rivers and perennial medium rivers, respectively, 7 present similar hydrographs with high base flow and moderate peak flows in February 8 or April and minimum flows in July or August. Differences on flow magnitude between 9 these classes are due to their environmental characteristics (Fig. 7), defined by their location in the Segura Basin (Fig. 8). Class 1 (MADIS > 10  $\text{m}^3/\text{s}$ ) includes medium and 10 11 low sections of the Segura River (Strahler order 5) with large drainage areas (more than 12 5000 km<sup>2</sup>), medium altitude (around 800 m.a.s.l) and slope (around 20%) and an annual mean precipitation of 450 mm. However, Class 2 (MADIS =  $2-10 \text{ m}^3/\text{s}$ ) corresponds to 13 14 upper sections (Strahler order 3) of the Segura River as well as medium and low 15 sections of the Mundo River, in wetter (700 mm of average precipitation) and highly 16 karstified (75% mean karstic surface) watersheds. These watersheds are higher than 1100 m.a.s.l., smaller than 2000  $\text{km}^2$  and have a 30% of slope. 17

18

19 The rest of hydrological classes, tributaries, follow environmental gradients (Fig. 7). 20 Classes 3 (*perennial creeks*) and 4 (*perennial headwater streams*) correspond to 21 headwater streams dominantly of orders 2 and 1, respectively, located in the upper 22 sectors of the Segura Basin with an average karstic surface in their watersheds greater 23 than 70%. These classes are characterized by soft (groundwater-driven) hydrographs 24 with flows varying among streams for most months but higher in winter than in summer 25 (Fig. 6). However, classes 5 (*seasonal winter-spring streams*) and 6 (*seasonal spring*)

1 streams) comprise streams with similar flows during summer-autumn but different in 2 winter-spring. They present maximum flows in December and March (Class 5) or only in March (Class 6) due to seasonal precipitation peaks. For these classes, watersheds 3 4 were low (less than 40%) and medium (around 50%) karstified respectively. Class 5 5 includes medium size streams (orders 3-4) that rarely dry up, located principally in the 6 medium (800 m.a.s.l) elevations of the Segura Basin. Class 6 is composed of springs 7 located in the headwaters of small watersheds with similar altitude and slightly higher 8 slope (Fig. 7), in any sector of the basin, that can cease flowing during less than one 9 month per year. However, streams in class 5 presented higher variability in annual flows 10 than streams in class 6 (Fig. 5).

11

12 Classes 7 and 8 (intermittent streams and ephemeral streams, respectively) have the 13 smallest mean annual flows, but the largest coefficients of variation for both annual 14 (Fig. 5) and monthly flows. They are characterized by intense and frequent droughts and 15 flash floods. Intermittent streams presented more predictable flows (Fig. 5) and softer 16 peaks (Fig. 6) than ephemeral streams. Associated to strong rain events, these peaks are 17 punctual in spring (March) and sustained in autumn (October-November). However, 18 ephemeral streams presented a higher coefficient of variation (Fig. 5) and only a peak of 19 flow (Fig. 6) in winter (December), greater than the ones for intermittent rivers. This 20 peak is associated to torrential precipitation episodes that compose most annual water 21 resources in this class. Both intermittent and ephemeral streams present low orders (1-2) and small drainage areas (less than 150 km<sup>2</sup>), restricted to the southern half of the 22 23 Segura Basin, in areas of low altitude (around 600 m.a.s.l), small slope (around 15%), 24 reduced karstic surface (close to 30% and 5%, respectively) and low average 25 precipitations (Fig. 7).

1

# 2 **Discussion**

3

From the 73 hydrological indices studied, three metrics describe the patterns of hydrological variability in the Segura River Basin: mean annual flow (MADIS), interannual coefficient of variation ( $CV_{INTER}$ ) and duration of droughts ( $D_L$ ); since they represent the dominant gradients detected on flows: (1) magnitude and (2) temporal variability. They reflect the specific hydroclimatic characteristics of the study region: scarce and irregular precipitation as well as discharge associated with hydrological extremes (drought and floods), typical for Mediterranean areas (Gasith and Resh 1999).

11

The ability to reduce the population of indices to a small, manageable subset has a number of benefits, including the reduction of analytical time and resources (Olden and Poff 2003). Other classification studies in Mediterranean rivers have used similar hydrological variables related to flow magnitude, variability and drought intensity as the main discriminators of flow regime classes (Baeza and others 2006), sometimes in combination with morphological, geological and climatic variables following the system B of the EU WDF (Munne and Prat 2004; Sanchez-Montoya and others 2007).

19

The distribution of flow-regime classes showed a high degree of spatial cohesion (Fig. 8), with most classes following the aridity gradient from NW to SE in the Segura River Basin. The most permanent and regular flows were found in the NW and the most intermittent and irregular flows in the SE. This regular-irregular flow gradient found in the Segura River Basin is similar to the observed by Baeza and others (2006) in the Tagus River Basin in central Spain, Poff and Allan (1995) in the rivers of Wisconsin

and Minessota and Mcnamay and others (2011) in the southeastern US. However, the
 flow regime Class 6, described as *seasonal spring streams*, is broadly distributed across
 all sectors of the Segura Basin in small and scarce karstic watersheds.

4

5 Differences on geology along the NW-SE gradient, coupled with climatic differences, 6 explain the differences of base flow among hydrological classes. The upper sector (high 7 elevations) of the Segura Basin is more karstic than the medium and low sectors, which 8 determines a more stable and regular hydrograph in the classes 3 and 4 (perennial 9 creeks and perennial headwater streams, respectively), located in the Northwest. 10 However, in the opposite extreme, the dominance of impermeable sediments (clay and 11 marls) produces quick runoff and flashy hydrographs, characteristic of classes 7 and 8 12 (*intermittent* and *ephemeral streams*, respectively).

13

14 With more than a third of all the nodes and a drainage area greater than 60%, 15 intermittent and ephemeral streams are the predominant classes in the Segura Basin, as 16 in other arid and semiarid areas of Australia (Boulton and Suter 1986) and South Africa 17 (Davies and others 1993; Uys and O'Keeffe 1997). In these streams high flow variability 18 indicates periods without flows, whereas in perennial streams it denotes fluctuations 19 (Uvs and O'Keeffe 1997), making difficult to establish discrete classes along the 20 temporal variability continuum. However, the duration and periodicity of no-flow 21 phases, the season when flow peaks occur and the variability in flow regimes within and 22 among years are key components to define and characterize these streams.

23

In other Mediterranean basins, like the Ebro Basin, the duration and timing of low flowsare the most important hydrological variables to discriminate flow regime classes

1 (Bejarano and others, 2010). We considered the drought duration as the most important 2 parameter because it was correlated with all the studied metrics and represents the 3 gradient of temporal variability in the Segura basin. The drought duration metric has 4 ecological significance emphasizing the biological consequences of the intensity of 5 droughts (Martinez and Fernandez 2006). It is probably the most important 6 environmental parameter affecting the aquatic biota in temporary rivers (Boulton 1989). 7 Drought events can result in the stream channel drying, partially or completely, and 8 both aquatic space and quality declining, which undoubtedly affect organisms. Drought 9 play a key role in the distribution of species, community structure and life-history 10 strategies of resident species (Gasith and Resh 1999), although some responses are 11 stream and community-specific (Argerich and others 2004; Dewson and others 2007). 12 Although droughts in Mediterranean climatic regions are predictable and periodic 13 (Gasith and Resh 1999), their intensity can vary because of interannual variations in 14 weather (Boix and others 2010). In Mediterranean climates, native biota have life 15 history traits that provide them with greater resistance to droughts and an improved 16 ability to get over a disturbance (Bonada and others 2007; Ferreira and others 2007), but 17 may make them particularly vulnerable to the alteration of flow regimes (Lytle and Poff 18 2004).

19

Human activities both in streams (e.g., flow regulation) and catchments (e.g. agriculture and urbanization) can exacerbate droughts and floods (Lake 2007), especially in Mediterranean areas densely populated with intense water abstraction and regulation. In the Segura River, and some tributaries, reservoirs profoundly alter the natural flow regime, causing a significant reduction in the magnitude of flows and a relevant modification of the seasonal pattern, with droughts during winter (instead of summer

1 months) becoming more frequent and durable (Belmar and others 2010; Vidal-Abarca 2 and others 2002). The effects of these alterations on ecosystem structure and 3 functioning are poorly known in the basin. In other Mediterranean rivers, Boix and 4 others (2010) found that reservoirs intensified the effect of droughts on the composition 5 and structure of diatoms and fish assemblages downstream of dams. Besides, the 6 decrease of flood frequency and the occurrence of extended droughts facilitate the 7 invasion of exotic species, as occur in other regulated rivers (Lake 2003).

8

9 The hydrological classification scheme obtained provides a first level mean of 10 arranging, conceptualizing and describing the natural or "reference" flow regimes in the 11 study area at two levels of resolution. Despite the absence of components related to the 12 frequency, duration and rate of change of high flow events, due to the use of monthly 13 data, a functional classification was obtained. Like in other hydrological classifications 14 that used monthly flow records (Harris and others 2000, Bejarano and others 2010), or 15 even daily gauged data (Mcnamay and others 2011), important spatial and temporal 16 variations in hydrologic characteristics were detected. Therefore, monthly data may be 17 adequate to analyze peak flows in Mediterranean streams, characterized by their 18 seasonal maximums, and this classification is potentially relevant to develop 19 environmental flows in the study area considering the magnitudes of the high flows 20 necessary for an environmental regime (Poff, 1996). Similarly, monthly flows may be 21 useful to determine the magnitude and duration of low flow events, which generally 22 present larger duration than high flow events. However, monthly flows present some 23 limitations to the design of environmental flows, such as the determination of the rise 24 and fall rates during extreme events, which require daily or hourly flow series (Bejarano 25 and others 2010).

1

2 The resulting classification will provide a strong basis for the study of the flow 3 alteration-ecological response relationship in each hydrological type, a critical step to 4 assess environmental flows within the ELOHA framework (Poff and others, 2010). The 5 comparison between the obtained reference flows and the actual ones, determined from 6 gauging data, will allow us to characterise the hydrological alteration in each river type. 7 Then, the flow alteration-ecological response relationship will be established by 8 biological monitoring in sites selected along the gradient of hydrologic alteration. The 9 development of this relationship for different river types will provide flow standards for 10 water managers to guide the development of environmental flows both for rivers and for 11 river segments in the Segura Basin.

12

In summary, the resulting classification is an example of a reference hydrologic classification in a Mediterranean basin where there are very limited unaltered flow data and only modelled monthly flows are available. A useful tool to support ecologically sustainable water resources planning and management in the Segura River Basin within the ELOHA framework.

18

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20

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Identification code	ΝΤ	Hydrologic index	Description	Units	
Magnitude of flow events					
Average flow conditions					
M <sub>A</sub> 1 - 12	12 M	Mean monthly flows	Mean monthly flow for all months	m³⋅s⁻¹	
CV <sub>A</sub> 1 - 12	12 M	Variability in monthly flows	Coefficient of variation in monthly flows for all months	-	
M <sub>A</sub> 13 - 14	2 M	Variability across monthly flows 1	Variability in monthly flows divided by median monthly flows,	-	
			where variability is calculated as range and interquartile		
CVINTRA	1 M	Variability across monthly flows 2	Coefficient of variation in mean monthly flows		
M <sub>A</sub> 15	1 M	Skewness in monthly flows	(Mean monthly flow-median monthly flow)/median monthly flow	-	
M <sub>A</sub> 16	1 A	Mean annual runoff	Mean annual flow divided by catchment area	-	
M <sub>A</sub> 17 - 18	2 A	Variability across annual flows	Variability in annual flows divided by median annual flows,	m <sup>3</sup> ⋅s <sup>-1</sup> ⋅km <sup>-2</sup>	
			where variability is calculated as range and interquartile	-	
M <sub>A</sub> 19	1 A	Skewness in annual flows	(Mean annual flow—median annual flow)/median annual flow	-	
Q1	1 A	Variability across annual flows 1	Percentile flow with the annual discharge exceeded	m³⋅s⁻¹	
	• •		1% of the time		
Q5/Q50, Q10/Q50	2 A	Variability across annual flows 2	Percentile flows with the annual discharge exceeded	-	
050	1Δ	Median annual discharge	Median annual flow for all years	m <sup>3</sup> .e <sup>-1</sup>	
MEDDIS/A	1 A	Median annual runoff	Median annual discharge divided by catchment area	m <sup>3</sup> .s <sup>-1</sup> .km <sup>-2</sup>	
RANGE	1 A	Range of flows	Maximum annual discharge minus minimum annual discharge	m <sup>3</sup> ·s <sup>-1</sup>	
STDEV	1 A	Variability across annual flows 3	Standard deviation of annual discharge	m <sup>3</sup> ⋅s <sup>-1</sup>	
CVINTER	1 A	Variability in annual flows	Coefficient of variation in annual flows for all years	-	
MADIS	1 A	Mean annual discharge	Mean annual flow for all years	m³⋅s⁻¹	
Low flow conditions		Ū.	·		
M <sub>L</sub> 1 - 12	12 M	Mean minimum monthly flows	Mean minimum monthly flow for all months	m³⋅s⁻¹	
M <sub>L</sub> 13	1 A	Average minimum monthly flow	Mean of the mean minimum flows for all months	m <sup>3</sup> ⋅s <sup>-1</sup>	
AMIN/Q50	1 A	Annual minimum	Minimum annual discharge divided by Q50	-	
I <sub>L</sub>	1 A	Drought intensity	Monthly flow equalled or exceeded 95% of the time divided	-	
			by mean annual flow		
High flow conditions				2 1	
M <sub>H</sub> 1 - 12	12 M	Mean maximum monthly flows	Mean of the maximum monthly flows for all months	m³⋅s⁻¹	
CV <sub>H</sub>	1 M	Variability across maximum monthly flows	Coefficient of variation in mean maximum monthly flows	-	
M <sub>H</sub> 13	1 M	Average maximum monthly flow	Mean of the mean maximum flows for all months	m³⋅s⁻¹	
AMAX/Q50	1 A	Annual maximum	Maximum annual discharge divided by Q50	-	
IH	1 M	Flood Intensity	Monthly flow equalled or exceeded 5% of the time divided	-	
by mean monthly flow					
	1 M	Percent of zero-flow months	Percentage of all months with zero flow	%	
<u> </u>	1 171		r oroontage of all months with 2010 now	70	

Appendix: Hydrological indices used for hydrological classification. (N: number of indices; T: time basis, being "M" monthly and "A" annual).

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#### Captions for figures

Figure 1. Location of the Segura River basin in Spain, showing the drainage network and nodes (black points) obtained from a digital elevation model

Figure 2. Two-dimensional PCA ordination of the 390 stream and river sections showing the correlated hydrological metrics (see Appendix for definition), the gradients detected (magnitude and temporal variability) and the hydrological class for each stream according to the clustering in Figure 3

Figure 3. Dendrogram obtained of the flexible- $\beta$  clustering procedure carried out with Euclidean distances. Two levels of classification, eight and four hydrological classes (see dotted lines), and the critical values of hydrological metrics that best discriminate them are showed

Figure 4. Evolution of the ANOSIM R-Value as the number of classes resulting from the flexible- $\beta$  clustering increases

Figure 5. Box-plots for the comparison of the duration of droughts ( $D_L$ ), the coefficient of interannual variation ( $CV_{INTER}$ ) and the magnitude of annual flows (MADIS) for the eight hydrological classes defined in the Segura River Basin. Names of classes detailed in Figure 3

Figure 6. 90th and 10th percentiles (in bars) and means (solid circles) of standardized monthly flows (monthly flows divided by its median) of all river and stream sections

included in each hydrological class. Ordinates are showed at different scales to improve the visualization

Figure 7. Box-plots for the comparison of the environmental variables (average precipitation in the drainage area, drainage area, Strahler order, average altitude as well as slope of the drainage area and percentage of karstic surface) among the eight hydrological classes defined in the Segura River Basin. Names of classes detailed in Figure 3

Figure 8. Map of the Segura River Basin showing the river segments and the 8 hydrological classes defined by Euclidean distances flexible- $\beta$  clustering





- ▼ Class 4: Perennial headwater streams
- Class 5: Seasonal winter-spring streams
- ▲ Class 6: Seasonal spring streams
- + Class 7: Intermittent streams
- \* Class 8: Ephemeral streams











Min-Max

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