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Additional Information

# 1 Thermal response test analysis for U-pipe vertical borehole heat 2 exchangers under groundwater flow conditions

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## 12 ABSTRACT

13  
14  
15 Conventional models used in the analysis of thermal response test data only consider conduction as heat  
16 transfer mechanism. In cases where presence of groundwater is detected, convection heat transmission  
17 plays an important role, so its influence must be determined in the calculation of the effective thermal  
18 conductivity, usually overestimated in these situations, increasing its value the higher the power injected  
19 and the time elapsed. In this work, based on the data collected in a borehole located at UPV (València)  
20 in which have been carried out three thermal response tests with different characteristics, has been  
21 implemented a variation of the finite line source model introducing an expression for the effective  
22 thermal conductivity formed by two terms, one static unaffected by underground flow and another  
23 dynamic that depends on time. Analyzing the data in the model developed and in the finite line source  
24 and infinite line source models, the results show that the new model estimates accurately the  
25 conductivity value unaffected by underground flow regardless the power injected or the time elapsed in  
26 the test, with differences between the results obtained in the analyzed tests and average thermal  
27 conductivity of 1,4%, compared to the conventional models in which this difference is 27%.

28  
29 **Keywords:** Thermal response test (TRT) analysis; Geothermal heat exchanger; Ground water advection; Effective  
30 thermal conductivity; Borehole thermal resistance; Undisturbed ground temperature recovery

31

## NOMENCLATURE

32		
33		
34		
35	$\alpha$	<i>ground thermal diffusivity</i>
36	$C_v$	<i>ground volumetric thermal capacity</i>
37	$\gamma$	<i>Euler constant</i>
38	$Ei$	<i>Euler integral</i>
39	$\lambda$	<i>effective ground thermal conductivity</i>
40	$\lambda_0$	<i>true ground thermal conductivity unaffected by groundwater flow</i>
41	$L$	<i>borehole depth</i>
42	$\dot{m}$	<i>fluid mass flow</i>
43	$Q_z$	<i>constant heat power injected to the ground per length unit</i>
44	$R_b$	<i>borehole thermal resistance</i>
45	$r_b$	<i>borehole radius</i>
46	$T_0$	<i>undisturbed ground temperature</i>
47	$T_{in}$	<i>fluid inlet temperature</i>
48	$T_{out}$	<i>fluid outlet temperature</i>
49	$T_{ave}$	<i>average of the fluid temperature</i>
50	$T_b$	<i>temperature at the borehole surface</i>
51		
52		
53		

## 54 1. INTRODUCTION

55 The most commonly used method to obtain the necessary data for ground source heat  
56 exchangers (GSHE) proper design in medium or large installations is the thermal response test  
57 (TRT), a procedure technically and economically accepted by designers [1-2] and promoters of  
58 shallow geothermal facilities, being used for more than two decades [3-4]. Thermal values  
59 obtained by following the indications described in regulations and standards [5-7] usually have  
60 no discussion among GSHE designers, regardless the limitation of the application of the model,  
61 the different ground characteristics and the measurements conditions in the work site. However,  
62 the measurement and analysis of the thermal ground parameters: ground thermal conductivity,  
63 borehole thermal resistance and undisturbed ground temperature [8-9] can be conditioned for  
64 different reasons. Actually, thermal conductivity measured in a thermal response test is called  
65 effective thermal conductivity because, due to the effects of an inhomogeneous ground and  
66 possible presence of groundwater flow, the heat transfer process is not pure conductive.

67 Thermal response test is carried out connecting a mobile equipment formed by a heating  
68 or cooling system, a hydraulic pump, flow and temperature sensors and a control system to a  
69 geothermal probe installed in a borehole in order to inject or extract a constant thermal power.  
70 Therefore, the first aspect to take into account in a TRT is the equipment control system, which  
71 must guarantee to perform the test under constant power conditions. For example, using a PID  
72 control system allows a more accurate analysis by reducing the error associated to the  
73 measurements [10]. Secondly, considering that the main outputs of the TRT are the inlet and  
74 outlet temperature of the heat carrier fluid as a function of time, minimizing the length of the  
75 connection pipes between the TRT equipment and the borehole should be a priority although  
76 sometimes it is not possible due to the work site conditions. In these cases, a filtering technique  
77 of the undesired effect produced in fluid temperature measurements by the ambient temperature  
78 can be used [11].

79 For data analysis and thermal parameters characterization different models are used  
80 [12], the most widely applied method is the infinite line source [13–17] but other approaches  
81 such as the finite line source model [18–21] or cylindrical source model [22–25] are also well  
82 known. These analytical models are used because of their simplicity and good accuracy of the  
83 results, mainly the infinite line source model, but a limitation to this methodology is the amount  
84 of groundwater flow [26]. As the effective ground thermal conductivity determined in TRT  
85 includes convection effects, in these cases its value is strongly conditioned. Advective  
86 phenomena, that is how groundwater flow transport the heat injected what depends on  
87 groundwater velocity, which is related to the hydrogeological characteristics of the different  
88 ground layers [27], are not being considering in the heat transfer models mentioned.

89 The effects of groundwater natural convection on borehole thermal resistance have been  
90 studied in groundwater-filled boreholes [28-29]. In grouted boreholes installed under  
91 groundwater flow conditions, the advective phenomena is relevant in borehole heat transfer  
92 boundary conditions. To consider it, several works proposed an analytical solution based on a  
93 moving finite line source model to consider the groundwater flow [30-33] introducing the Péclet  
94 dimensionless number in the heat transfer models. Other authors [34] have developed a new  
95 test protocol to evaluate the effects of convection and lateral groundwater flow based on the  
96 application of several heat injection and extraction pulses using a numerical model with a  
97 parameter estimation technique to obtain the thermal ground parameters. The incorporation of  
98 these proposals to the thermal response test analysis requires a computational effort, a longer  
99 period of data collection to minimize errors [35] or a reversible (heating and cooling) TRT  
100 equipment which supposes a barrier for the methodology standardization.

101           Although there are numerous methods to calculate the borehole thermal resistance [36]  
102 its estimation in situ by means of a TRT is important not only to obtain a value to carry out the  
103 shallow geothermal system dimensioning but also to verify the correct GSHE execution by  
104 comparing the measured with the expected value. Considering that the borehole resistance error  
105 is mainly influenced by the error of thermal conductivity [37], in ground source heat exchangers  
106 working under groundwater conditions, borehole resistance will also present the same positive  
107 effect than the effective thermal conductivity [38]. Performing several TRT varying the  
108 injection parameters [39–40] or enlarge the thermal response test duration [41] are adequate  
109 procedures to characterize properly the borehole thermal resistance. In this work, an accurate  
110 analysis of this parameter using three different heat injection ratios during long time periods is  
111 done.

112           Regarding the last ground thermal parameter that a TRT evaluates, undisturbed ground  
113 temperature, it is noted that this work is based on an exhaustive ground temperature  
114 characterization, measuring this value along the borehole depth before and during the thermal  
115 response test and evaluating the ground thermal recovery a long time after the end of the test,  
116 as is presented in previous works and collect in this contribution [42–43]. In this previous  
117 research, a standard analysis of thermal response test using infinite linear source model  
118 prediction can be found observing a deviation from the prediction of the infinite line model,  
119 showing an increase of the ground thermal conductivity with the advance of the test. To  
120 understand this phenomenon, the work developed in [44] presents an analysis procedure  
121 implemented by a 3D finite element model that completes the standard TRT analysis, estimating  
122 the thermal conductivity profile from a temperature profile along the borehole during the test.  
123 A highly conductivity layer was detected using this procedure, indicating the presence of  
124 ground water currents. The application of this methodology requires an extra resource effort  
125 regarding the TRT standard methodology because is needed a measurement of the borehole  
126 temperature profile and a more complex data analysis implemented in COMSOL. It should be  
127 noted that, in these previous works, the whole data recorded in during the research are not  
128 analysed; in this paper the complete data set are published.

129           The innovation introduced in this work is the presentation of a simple analysis  
130 methodology for standard thermal response tests performed under groundwater flow conditions,  
131 intended for engineering application, and based on finite line source model. This methodology  
132 is based on a phenomenological characterization of the impact of ground water advection in the  
133 estimation of ground thermal parameters extracted from line source model. The objective is that  
134 GSHE designers can know how the underground flow masks the result of the effective thermal  
135 conductivity and assess, based on this knowledge, what is the value that they will use for  
136 dimensioning, without the need for additional measurements to those made in standard TRT.

137           The structure of the paper is as follows: firstly, a description of the installation and the  
138 data collection system is done. Secondly, the characteristics of the three tests performed are  
139 presented and the raw data collected are analysed. Thirdly, an analysis of the effective thermal  
140 conductivity and the borehole resistance is carried out using traditional methodologies (finite  
141 line source and infinite line source analysis). Then, a modification of finite line analysis model  
142 to quantify the groundwater effects is presented and the data analysis redone using this new  
143 methodology. Finally, discussion of results and conclusions.  
144

145 **2. EXPERIMENTAL PLANT DESCRIPTION**

146 2.1 Site geological information and characteristics of the borehole heat exchanger

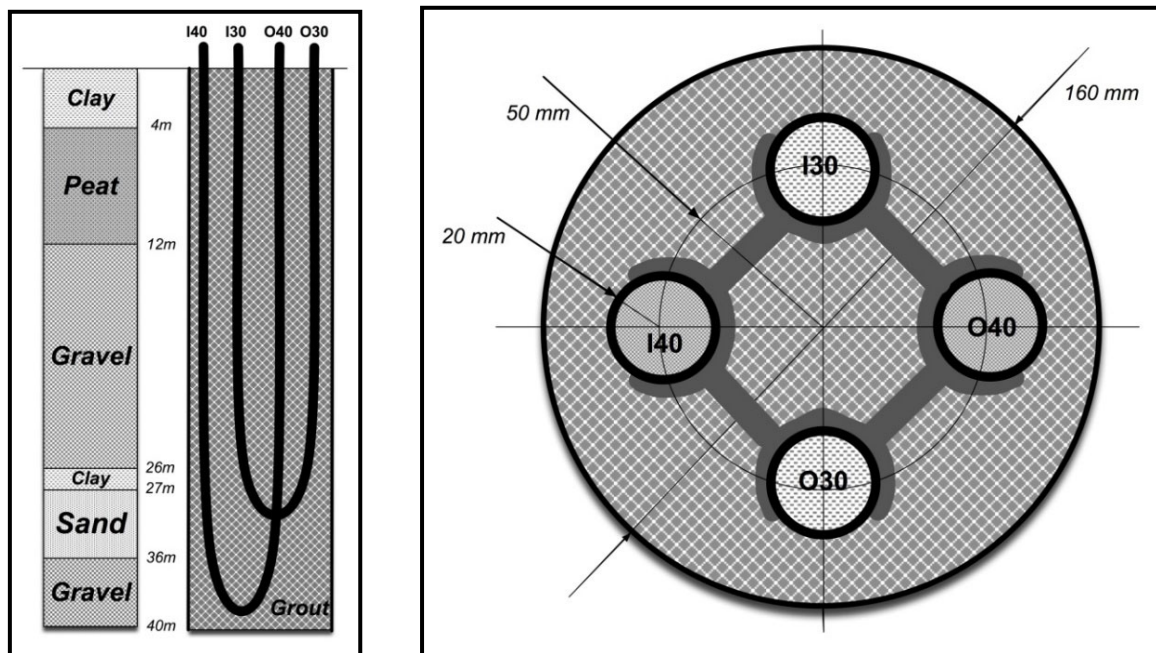
147

148 The experimental installation is located at Universitat Politècnica de València and it was built  
149 on the first days of May 2010. The test site presents geological characteristics representatives  
150 of Valencia city with gravels, sands and clays as predominant materials and a high groundwater  
151 flow presence. During the drilling works, six layers were identified along the 40 meters drilled,  
152 as can be seen at the stratigraphic column represented in figure 1, a clay layer from 0 to 4  
153 meters, a peat layer from 4 to 12 meters, a gravel with small round stones layer from 12 to 26  
154 meters, another clay layer from 26 to 27 meter, a sand layer from 27 to 36 meters and a last  
155 layer of gravel with small round stones from 36 to 40 meters.

156 The facility consists of a borehole of 40 m. depth in which two independent PE-100 U-  
157 pipes of 40 mm diameter were introduced. Initially was planned both pipes with 40 m. depth,  
158 but after executing the drilling inserting a non-extracting metallic casing, a narrowing in the  
159 initial borehole diameter (160 mm) was observed at 30 m. depth, due to the casing installation,  
160 so it was decided to introduce a shorter U-pipe (installed depth 29 m.) and a longer U-pipe of  
161 39 m. installed depth (figure 1). The space between the pipes and the casing was filled with a  
162 mixture of one part of bentonite and twelve parts of cement (CEMEX 32.5 raff) what is a very  
163 common commercial solution.

164 The installation is completed with a fixed thermo-hydraulic system that allows to carry  
165 out the heat injections test composed by a heating resistor of 3x1 kW/220 V, an electronic  
166 adjustable circulation pump and a 5-litre expansion tank (figure 2).

167



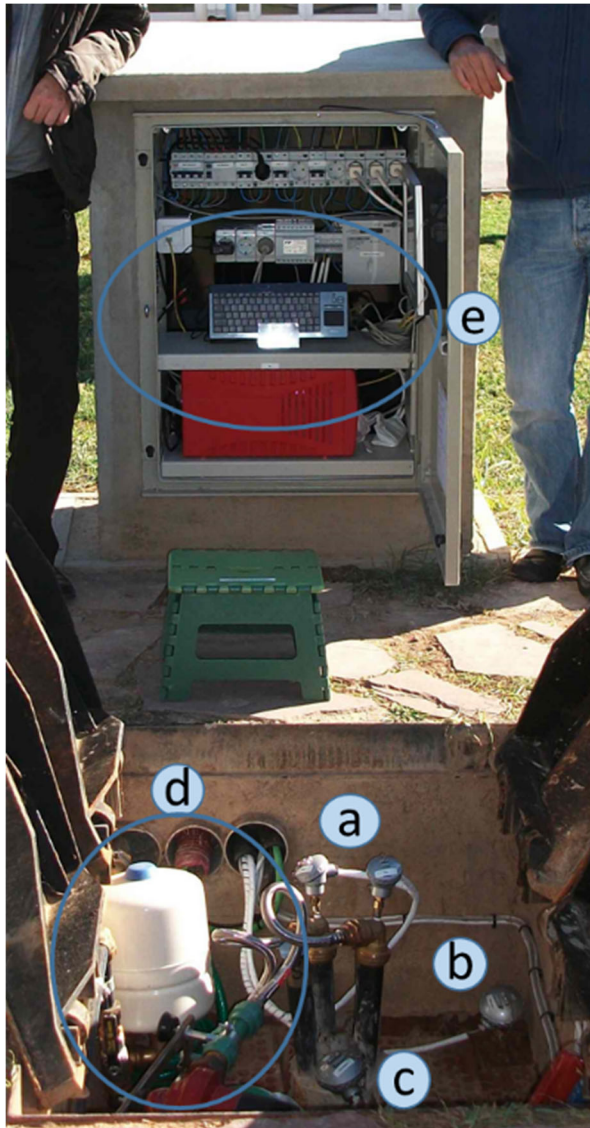
168 Fig. 1.- On the left, diagram showing the stratigraphic column of the borehole. On the right, diagram showing the location of  
169 both independent U-pipes, one 40 meters depth (I40, O40) and the other one 30 meters depths (I30, O30).

170

171 2.2 Monitoring system description

172

173 An equipment to control heat injection test was provided to the facility consisting of a flowmeter  
174 (accuracy of 1%) and temperature probes PT100 at input and output of the exchanger,  
175 connected to an acquisition system through a 4-wire 4–20mA loop of TC direct adjusted in a  
176 range from 0°C to 50°C. The temperature sensors (accuracy of 0,1 °C) were calibrated through  
177 a thermal bath and an electronic precision thermometer. Furthermore, an energy meter was  
178 employed for monitoring electric power source of the installation. The full system was managed  
179 from a PC with a touch screen and Internet access that performed acquisition and register of the  
180 data during the tests (figure 2).  
181



- a)  $T_{in}$ ,  $T_{out}$  temperature probes
- b) Ground temperature probe
- c) Borehole temperature probe
- d) Hydraulic subsystem
- e) Control and acquisition equipment

182

183 Fig. 2.- A picture of the borehole facility observing in the first place the borehole and the fixed thermo-hydraulic system and,  
184 in the background, the cabinet in which the data acquisition system is located.

185

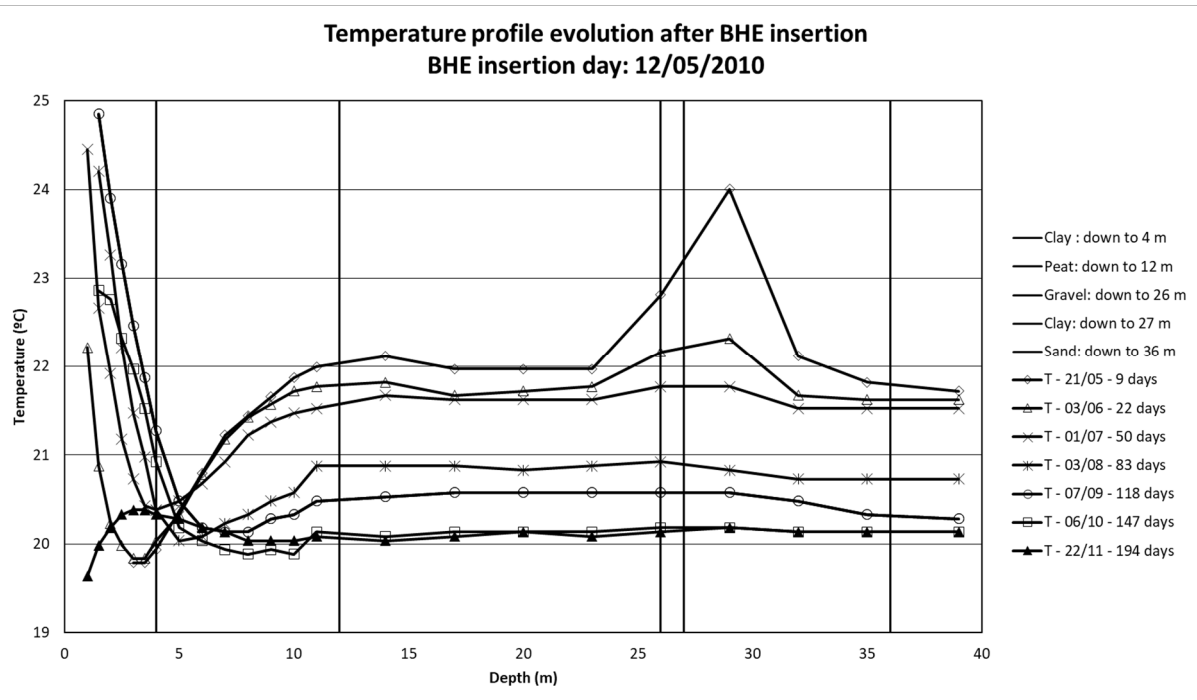
186 In order to regularly measure the ground temperature, the longer U-pipe installed was  
187 prepared as explained in [45].  
188

188

189 **3. THERMAL RESPONSE TEST MEASUREMENTS**

190 **3.1 Ground temperature characterization**

191  
192 During the six months after the borehole execution, the ground temperature was characterized  
193 by inserting a calibrate sensor in the longer U-pipe, measuring the water temperature inside the  
194 pipes, in thermal equilibrium with the surrounding ground. The measurement procedure,  
195 repeated at least once a month, consisted in descending the sensor at prefixed depths, holding  
196 it in the position for 5 seconds for thermal stabilization and moving to the next depth. Between  
197 1 and 4 meters depth, measures were taken every 0,5 meter, increasing that distance to 1 meter  
198 between 5 and 28 meters depth, reducing again the gap to 0,5 meters between 28,5 and 30,5  
199 because at this depth there were problems during the casing installation as explained, and  
200 ending measuring every meter between 30,5 to 39 meters. Figure 3 shows the temperature  
201 profile depending on depth obtained. It is noted that the average temperature decreases around  
202 2 °C during the monitoring period and the undisturbed ground temperature is reached the fifth  
203 month after the installation works. In the graph the different ground layers observed during the  
204 drilling have been marked with bold vertical lines, no significant changes in the temperature  
205 profile between them are observed but it is remarkable the temperature peak observed at 30 m  
206 depth. This effect is because during the drilling works at this depth a fracture in the casing was  
207 observed so the grouting spilled into the ground increasing its thickness at this depth. According  
208 to this, it can be concluded that the ground temperature behavior in the first months after drilling  
209 works is due to the heat released during the grouting setting. Through this analysis, it was  
210 determined that the undisturbed ground temperature ( $T_0$ ) at the test site is 20,12 °C.  
211



212  
213 Fig. 3.- Ground temperature profile as a function of depth. Measures are given for the following 2010 dates: 21/05, 03/06,  
214 01/07, 03/08, 07/09, 06/10 and 22/11.  
215

216 **3.2 Thermal response tests performed description**

217  
218 Once the undisturbed ground temperature was characterized, thermal response tests were  
219 started. Three heat injection tests were carried out in the shorter U-pipe during the following

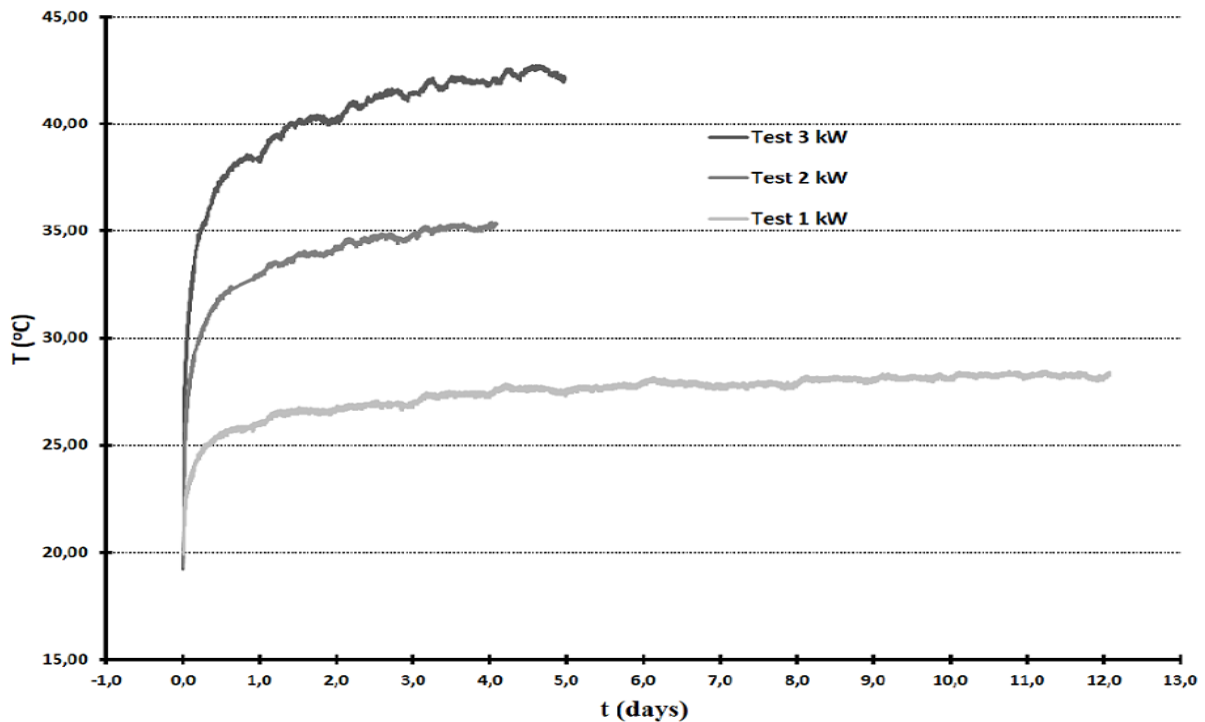


220 year and a half to the facility commissioning. The main characteristics of the test performed are  
 221 shown in table 1.  
 222

Test	Dates	Duration (hours)	Data acquisition period (seconds)	Average injected thermal power (W)
2 kW	22/11/2010 – 26/11/2010	97,9	180	1637±51
3 kW	15/12/2010 - 20/12/2010	119,4	180	2449±67
1 kW	09/03/2011 – 30/03/2011	289,7	30	798±39

223 Table 1.- Duration of the TRT performed and average thermal power injected.

224 Figure 4 shows the measured data of the average fluid temperature circulating through  
 225 the borehole heat exchanger for the three tests,  $T_{ave} = (T_{in} + T_{out})/2$ , as a function of time. In  
 226 the following analysis of the data, presented in next sections, each test will be labeled 1, 2 or 3  
 227 kW tests, meaning the value of the heating resistors used, different from the actual injected  
 228 thermal power. Data for the 1 kW test, with 21 days duration, will be analyzed only for the first  
 229 12 days.



230 Fig. 4.- Average fluid temperature, T, as a function of time for the three performed test.

231  
 232 3.3 Ground temperature profile and recovery analysis  
 233

234 As mentioned above, the longer U-pipe was used to measure ground temperature profile before,  
 235 during and after every test performed, inserting a calibrated temperature sensor in the pipeline,  
 236 and measuring the temperature of the ground along the geothermal probe. The period in which  
 237 ground temperature was recorded in each test is shown in table 2.

238

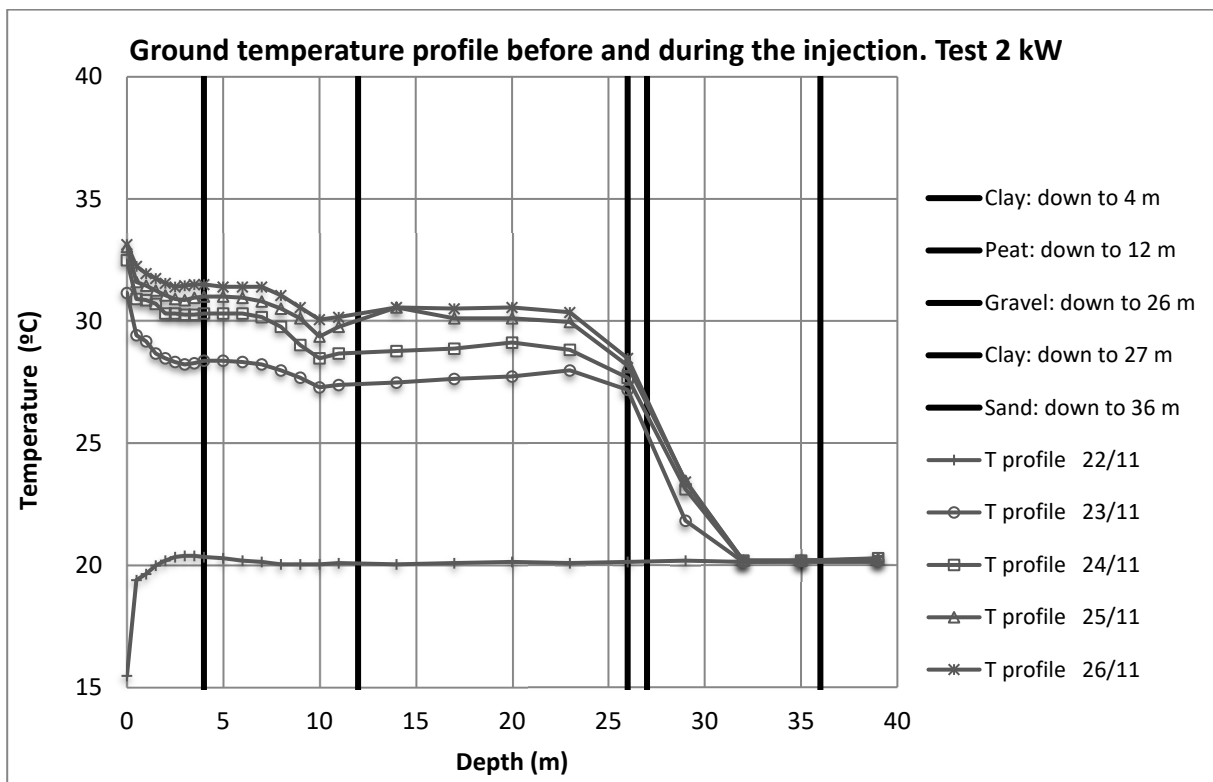
Test	Total period recorded	Duration of the recovery analysis period (days)
2 kW	22/11/2010 – 14/12/2010	18
3 kW	14/12/2010 – 10/01/2011	21
1 kW	02/03/2011 – 13/05/2011	44

239 Table 2.- Period of time in which ground temperature was recorded in each test.

240 Figures 5, 6 and 7 show the temperature values recorded before starting of the injection  
 241 and during the thermal test execution. In all test carried out it can be seen that, before the  
 242 injection, temperatures are quite constant, reaching the undisturbed ground temperature  
 243 estimated from 7 meters depth. A similar trend is also observed in the temperature profile during  
 244 the thermal test in the three power steps injected. Graphs present zones with higher heat  
 245 absorption capacity attributable to the presence of groundwater (around 10 meters deep, around  
 246 19 meters and around 25 meters). In figure 5 (2 kW test) is not observed in detail the temperature  
 247 decreasing at the depths of 19 and 25 meters because in this experiment the temperature  
 248 measurement was done each 3 meters between 10 and 39 meters. Due to the rapid drop of the  
 249 temperature observed between 25 and 30 meters deep and to perform a better analysis, in the  
 250 following tests (figures 6 and 7), the temperature was recorded every meter depth to observe  
 251 the behaviour in the different ground layers, clearly observing the zones with higher heat  
 252 absorption capacity.

253 The different behaviour between ground layers it is noted as well in the monitoring of  
 254 the ground several days after the completion of the thermal response test (figures 8, 9 and 10),  
 255 where same depths show a faster recovery. It is observed that the ground recovery period to  
 256 return to the undisturbed ground temperature is more than 15 days for all injected thermal power  
 257 values due to the long test period.

258



259 Fig. 5.- Ground temperature profile as a function of depth just before (22/11) and during the injection for the 2 kW test.  
 260  
 261

262  
263

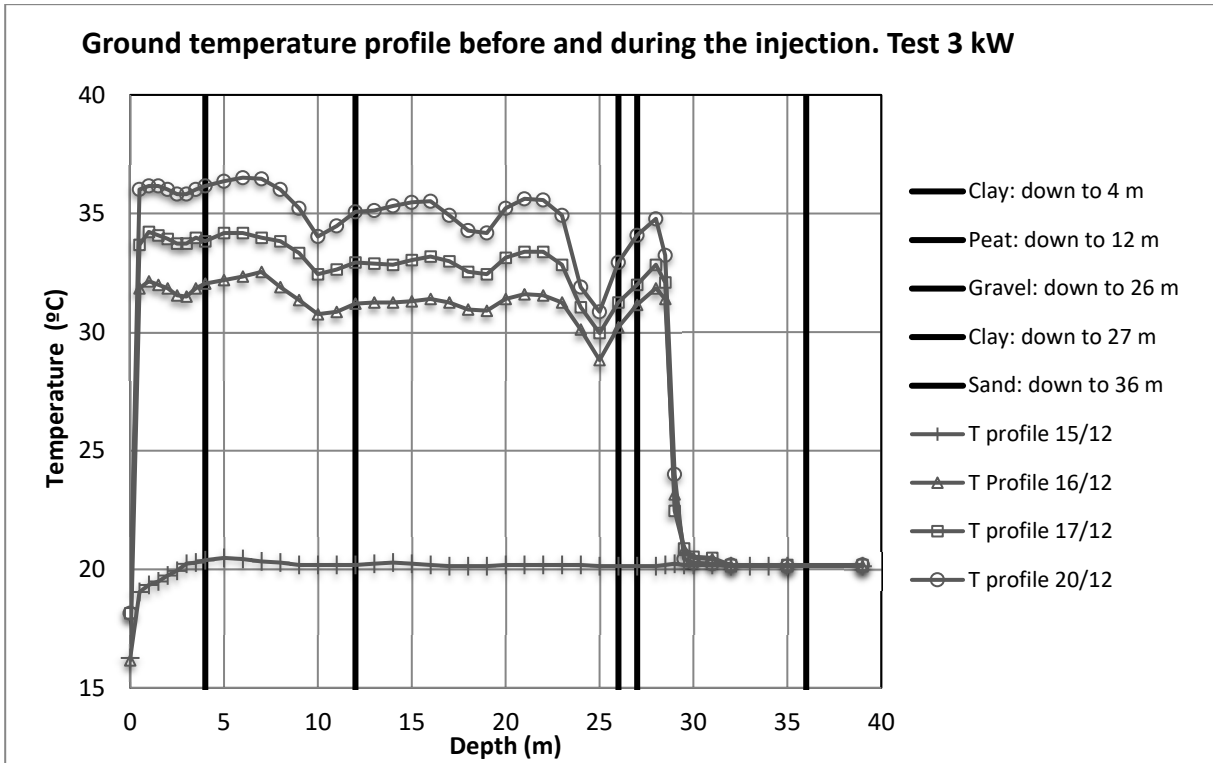


Fig. 6.- Ground temperature profile as a function of depth just before (15/12) and during the injection for the 3 kW test

264  
265

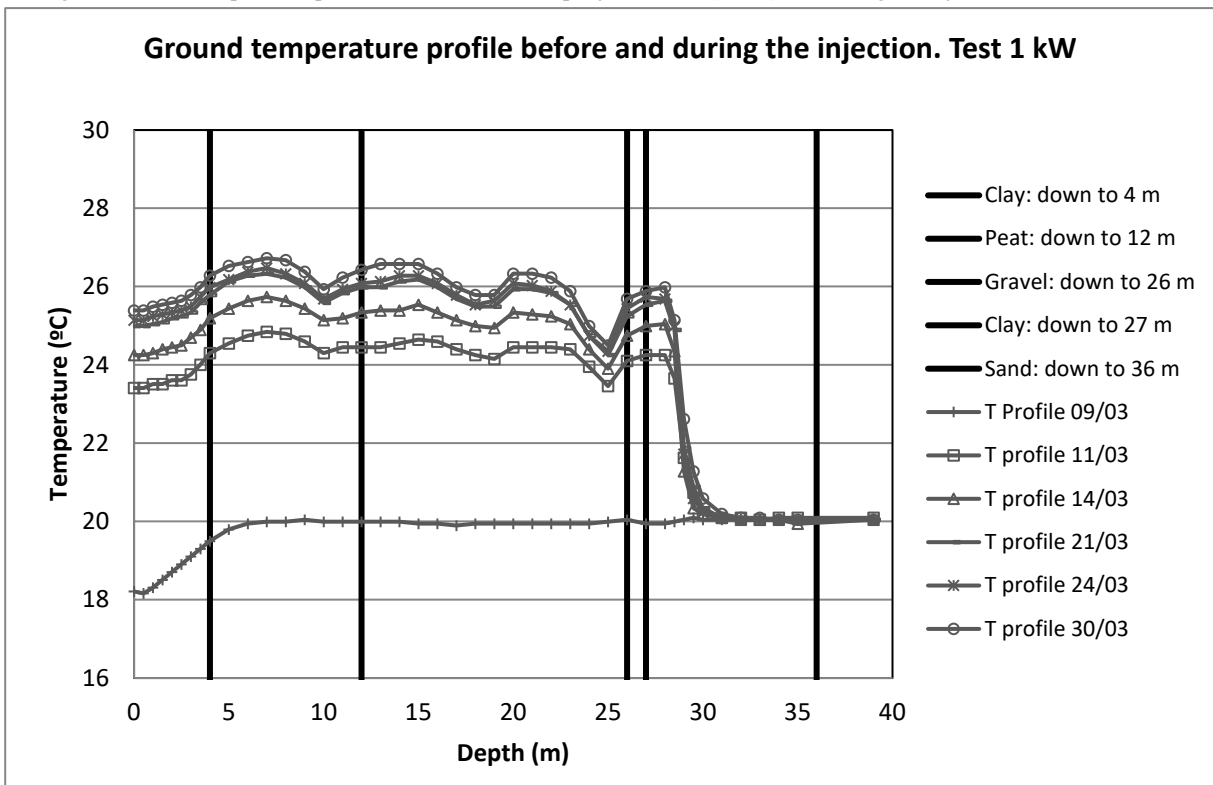
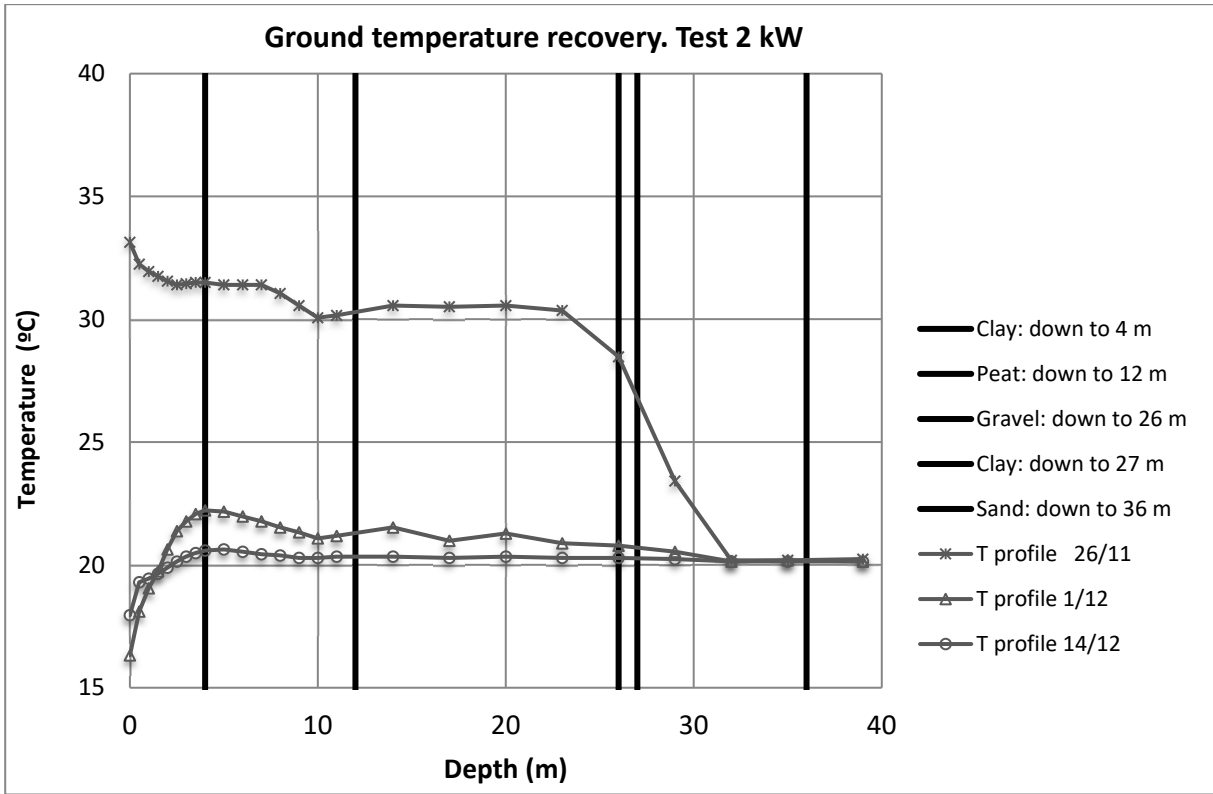
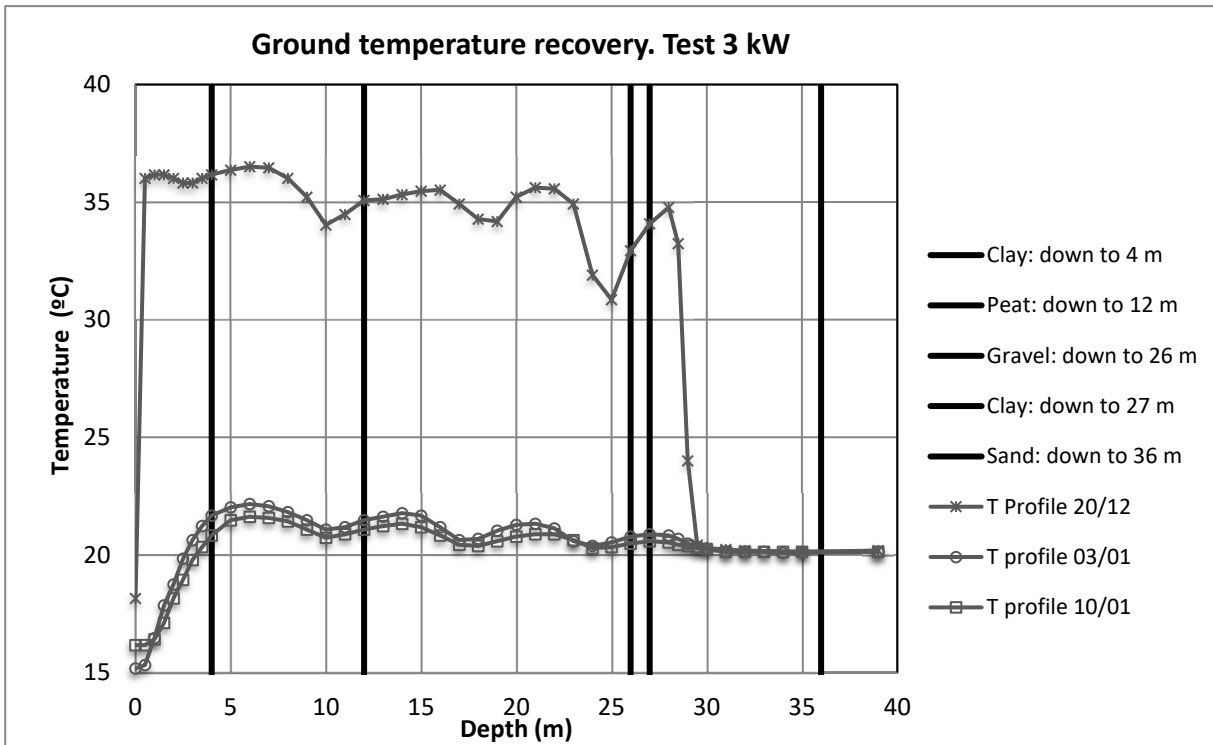


Fig. 7.- Ground temperature profile as a function of depth just before (09/03) and during the injection for the 1 kW test.



266  
267

Fig. 8.- Ground temperature profile as a function of depth at the end of the the injection (26/11) and after for the 2 kW test.



268  
269

Fig. 9.- Ground temperature profile as a function of depth at the end of the the injection (20/12) and after for the 3 kW test.

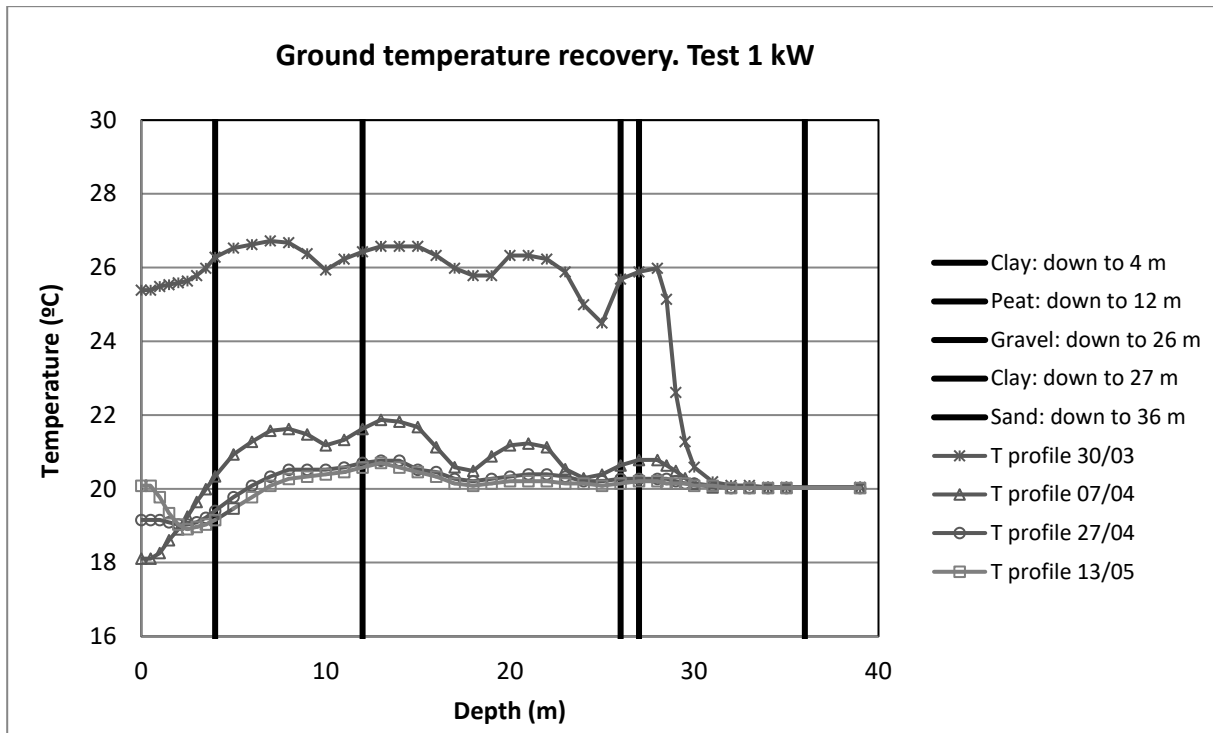


Fig. 10.- Ground temperature profile as a function of depth at the end of the the injection (30/03) and after for the 1 kW test.

270  
271  
272

273 **4. DATA ANALYSIS**

274 The purpose of this work is to obtain the effective ground thermal conductivity value under  
 275 groundwater conditions from data recorded on standard TRT, without the addition of new  
 276 measurements to characterize the subsoil layers, developing an effective model based on line  
 277 source analysis. This model is selected because is the most widely used in TRT analysis and is  
 278 recommended by international standards to find an accurate estimation of three parameters  
 279 approximately describing the thermal behaviour of the ground under consideration and needed  
 280 to design a ground coupled heat pump installation. These parameters are the undisturbed ground  
 281 temperature,  $T_0$ , the effective ground thermal conductivity,  $\lambda$ , and the borehole thermal  
 282 resistance,  $R_b$ . Variation of the line source model is carried out adding a new effective parameter  
 283 to incorporate the impact of ground water flow on ground thermal properties.

284 Line source analysis assumes that the borehole heat exchanger behaves as a linear heat  
 285 source emitting constant thermal power. This analysis also assumes that the ground is a  
 286 homogeneous infinite medium whose thermal behavior is characterized by its thermal  
 287 conductivity,  $\lambda$ , and its thermal diffusivity,  $\alpha$ . Considering the source with an infinite length,  
 288 meaning that the borehole depth,  $L$ , is much bigger than the borehole radius,  $r_b$ , the solution of  
 289 this thermal problem gives the temperature of the ground as a function of the radial coordinate,  
 290 and the time,  $t$ :

291 
$$T(r, t) = T_0 - \frac{Q_z}{4\pi\lambda} Ei\left(-\frac{r^2}{4\alpha t}\right) \quad (1)$$

292 Where  $T_0$  in the undisturbed ground temperature and  $Q_z$  the constant heat power injected  
 293 to the ground per length unit. Symbol  $Ei$  represents the Euler integral. For sufficiently large  
 294 times this expression can be approximated by:

295 
$$T(r, t) \approx T_0 + \frac{Q_z}{4\pi\lambda} \left\{ \ln \frac{4\alpha t}{r^2} - \gamma + \mathcal{O}\left(\frac{r^2}{4\alpha t}\right) \right\}, \quad for \quad \frac{4\alpha t}{r^2} \gg 1 \quad (2)$$

297 This expression is usually used to estimate the value of the temperature at the borehole  
 298 radius,  $r_b$ , during a Thermal Response Test:

299 
$$T(r_b, t) = T_b(t) = T_0 + \frac{Q_z}{4\pi\lambda} \left\{ \ln \left(\frac{t}{t_b}\right) - \gamma \right\}, \quad for \quad t \gg t_b = \frac{r_b^2}{4\alpha} \quad (3)$$

300 Then, borehole thermal resistance,  $R_b$ , is defined to model the inner problem of heat  
 301 transfer inside the BHE, relating the average of the fluid temperature,  $T_{ave}(t)$ , with the  
 302 temperature at the borehole surface,  $T_b(t)$ , through the expression:

303 
$$T_{ave}(t) = T_b(t) + Q_z R_b \quad (4)$$

304 Thermal response tests measure inlet,  $T_{in}$ , and outlet temperature,  $T_{out}$ , to the borehole  
 305 heat exchanger, as well as fluid mass flow,  $\dot{m}$ , allowing calculating average fluid temperature,  
 306  $T_{ave}$ , and thermal power injected to the ground,  $Q_z$ , through:

307 
$$T_{ave} = \frac{T_{in} + T_{out}}{2} \quad Q_z = \frac{\dot{m} c_p (T_{in} - T_{out})}{L} \quad (5)$$

308 If the assumptions of the infinite line source analysis are reasonable for the thermal  
 309 response test under consideration, then average fluid temperature will follow the expression:  
 310

311 
$$T_{ave}(t) = T_0 + Q_z R_b + \frac{Q_z}{4\pi\lambda} \left\{ \ln \left(\frac{t}{t_b}\right) - \gamma \right\} \quad (6)$$

312 Usual analysis plots data of average fluid temperature against logarithm of time. Then,  
 313 a linear behaviour of these experimental data will confirm the assumptions of the infinite line

314 source, extracting ground thermal properties and borehole resistance from the slope,  $a$ , and the  
 315 intercept,  $b$ , of the linear fit:

$$316 \quad a = \frac{Q_z}{4\pi\lambda} \quad b = T_0 + Q_z \left( R_b - \frac{\ln(t_b) - \gamma}{4\pi\lambda} \right) \quad (7)$$

317 From the slope,  $a$ , effective thermal conductivity of the ground is estimated,  $\lambda = Q_z/4\pi a$ ,  
 318 and from the intercept,  $b$ , a relationship between the undisturbed ground temperature,  $T_0$ , the  
 319 ground thermal diffusivity,  $\alpha$  (included in the time constant  $t_b = r_b^2/4\alpha$ ) and the borehole thermal  
 320 resistance,  $R_b$ , is found. If a measurement of the undisturbed ground temperature is done and  
 321 an estimation of ground thermal diffusivity is available, then borehole thermal resistance can  
 322 be calculated from expression:

$$323 \quad R_b = \frac{b - T_0}{Q_z} + \frac{\ln(t_b) - \gamma}{4\pi\lambda} \quad (8)$$

324 In the three thermal response tests analysed in this work an accurate measurement of the  
 325 undisturbed ground temperature was done and a previous estimation of ground thermal  
 326 diffusivity is available, so borehole thermal resistance can be calculated. As the three tests were  
 327 executed at different injected powers, it is convenient to choose appropriate variables allowing  
 328 comparing the three tests in a clear way, being a suitable choice  $f_0$  and  $\tau$ , defined as:

$$329 \quad f_0 = \frac{T_{ave}(t) - T_0}{Q_z} \quad \text{and} \quad \tau = \frac{\ln\left(\frac{t}{t_b}\right) - \gamma}{4\pi} \quad (9)$$

330 because the relation between these two variables predicted by the infinite line source approach,  
 331 given by:

$$332 \quad f_0 = R_b + \frac{\tau}{\lambda} \quad (10)$$

333 is independent of the injected power. Therefore, calculating variables  $f_0$  and  $\tau$  from the  
 334 experimental data of the three response tests, and plotting  $f_0$  against  $\tau$ , all experimental points  
 335 have to describe a line whose intercept is the borehole thermal resistance and its slope the  
 336 inverse of the effective thermal conductivity. First analysis of TRT data has been done using  
 337 this methodology. In addition, a second analysis with the purpose of evaluating the effect on  
 338 the estimates of a finite line source has also been done.

339 The solution of the heat transfer problem between the borehole heat exchanger and the  
 340 ground, considering a finite borehole length, is dependent of the vertical coordinate. Therefore,  
 341 the temperature at the borehole radius depends on this coordinate and the analysis of ground  
 342 thermal response test data is more elaborated. In this contribution, the procedure adopted to  
 343 consider finite length effects uses the average of the temperature at the borehole radius along  
 344 the whole length,  $L$ , of the borehole heat exchanger. The final expression to estimate the  
 345 temperature at the borehole radius is:

$$346 \quad T(r_b, t) = T_0 + \frac{Q_z}{4\pi\lambda} \left\{ \ln\left(\frac{t}{t_b}\right) - \gamma - \left( \frac{3}{\sqrt{\pi}} \left( \sqrt{\frac{t}{t_L}} - \left(\frac{r_b}{L}\right)^2 \sqrt{\frac{t_L}{t}} \right) - 3 \frac{r_b}{L} \right) \right\} \quad (11)$$

347 where  $t_L$  is a characteristic time scale associated to the borehole length:

$$348 \quad t_L = \frac{L^2}{4\alpha} \quad (12)$$

349 This expression for the temperature at the borehole radius was calculated in [20] as a  
 350 series expansion of the exact solution, averaged along the length of the BHE, in variables  $t_L$  and  
 351  $r_b/L$ . Note that if the length  $L$  is considered infinite the previous solution for the infinite  
 352 approach is recovered. Also, note the meaning of the infinite line approach, the length of the  
 353 borehole,  $L$ , is much bigger than the borehole radius,  $r_b$ .

354 Then, considering the relation between the average fluid temperature and temperature  
 355 at the borehole radius through the borehole thermal resistance, finite line source approach  
 356 predicts a temporal evolution of the average fluid temperature given by expression:

$$357 \quad T_{ave} = T_0 + Q_z R_b + \frac{Q_z}{4\pi\lambda} \left\{ \ln\left(\frac{t}{t_b}\right) - \gamma - \left( \frac{3}{\sqrt{\pi}} \left( \sqrt{\frac{t}{t_L}} - \left(\frac{r_b}{L}\right)^2 \sqrt{\frac{t_L}{t}} \right) - 3 \frac{r_b}{L} \right) \right\} \quad (13)$$

358 To analyse TRT data including finite lengths effects is also convenient a suitable choice  
 359 of variables to compare the three tests, being in this case  $f_0$  and  $\tau - \Delta\tau$ , with  $f_0$  and  $\tau$  as defined  
 360 previously and  $\Delta\tau$  defined as:

$$361 \quad \Delta\tau = \frac{1}{4\pi} \left\{ \frac{3}{\sqrt{\pi}} \left( \sqrt{\frac{t}{t_L}} - \left(\frac{r_b}{L}\right)^2 \sqrt{\frac{t_L}{t}} \right) - 3 \frac{r_b}{L} \right\} \quad (14)$$

362 because the relation between these  $f_0$  and  $\tau - \Delta\tau$  predicted by the line source approach, given by:

$$363 \quad f_0 = R_b + \frac{\tau - \Delta\tau}{\lambda} \quad (15)$$

364 is again independent of the injected power. Therefore, as with the infinite line approach, the  
 365 experimental values of  $f_0$  represented against  $\tau - \Delta\tau$  must describe a line whose intercept is the  
 366 borehole thermal resistance and its slope the inverse of the effective thermal conductivity.

367 Figure 11 is elaborated to show the accuracy of the finite line source prediction,  
 368 equations 13 and 15, to describe the behaviour of the experimental data acquired during the  
 369 execution of the three thermal response tests carried out. For calculating  $f_0$  (from equation 9)  
 370 variable the values of the undisturbed ground temperature,  $T_0$ , and the average injected thermal  
 371 power,  $Q_z$ , are needed. Undisturbed ground temperature has been accurately measured resulting  
 372 the value  $T_0=20,12$  °C. Average injected thermal power per length unit,  $Q_z$ , is calculated from  
 373 the experimental measurements of average injected power shown in table 1, resulting the values  
 374 presented in table 3. And for calculating  $\tau$  and  $\Delta\tau$  (from equations 9 and 14) the values of the  
 375 borehole radius ( $r_b=0,08$  m), borehole depth ( $L=29$  m) and the ground thermal diffusivity,  $\alpha$ ,  
 376 are needed. An estimation of the ground thermal diffusivity is available from previous research  
 377 works, from reference [46], in which a comparison between design and actual energy  
 378 performance of a HVAC-ground coupled heat pump system located 500 m. away from the test  
 379 site is presented, reporting values for ground thermal conductivity  $\lambda=1,43$  W/mK, ground  
 380 volumetric thermal capacity  $C_v=2400$  kJ/m<sup>3</sup>K and ground thermal diffusivity  
 381  $\alpha=\lambda/C_v=0,0000006$  m<sup>2</sup>s<sup>-1</sup>. This constant value of ground thermal diffusivity,  $\alpha$ , is the one used  
 382 along the present analysis.

383 Figure 11 represents the values of  $f_{0i}=f_0-R_{b0i}$  against the values of  $\tau - \Delta\tau$  calculated from  
 384 these experimental data, with  $R_{b0i}$  the borehole thermal resistance estimated for the test  $i$   
 385 ( $i=1,2,3$ ) at the beginning of each test. If FLS prediction applies, then all experimental points  
 386 will show a linear behaviour with 0 intercept and slope the inverse of the effective thermal  
 387 conductivity. The choice of variable  $f_{0i}$  for the vertical axis of figure 11 is done to enhance  
 388 clarity of the comparison between the three tests. A linear fitting between  $f_0$  and  $\tau - \Delta\tau$ , of the  
 389 data from values of  $\tau - \Delta\tau$  starting in 0,03 (equivalent to 1,9 hours) and finishing in 0,07  
 390 (equivalent to 3,2 hours) for the three tests, gives the values included in table 3 for the intercept  
 391 and the slope. Then,  $R_{b01}=0,094$  mK/W,  $R_{b02}=0,117$  mK/W and  $R_{b03}=0,114$  mK/W.

392

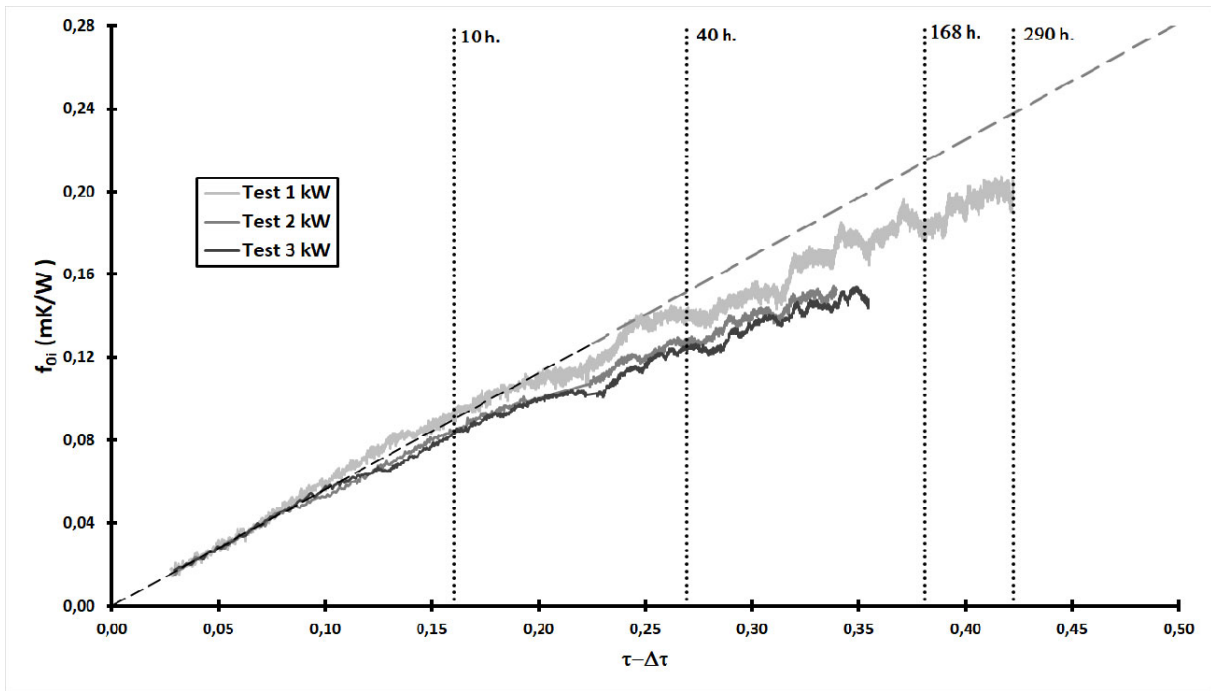
Test	Injected power $Q_z$ (W/m)	Intercept (m K/W)	Slope (m K/W)	C.C.
1 (1 kW)	27,72	0,094	0,564	0,974



2 (2 kW)	56,45	0,117	0,564	0,995
3 (3 kW)	84,44	0,114	0,560	0,994

393 Table 3.- Values of intercept, slope and correlation coefficient (C.C.) for the linear fitting of  $f_0$  against  $\tau-\Delta\tau$  for data of  $\tau-\Delta\tau$   
394 between 0,03 and 0,07.

395 A first estimation of the effective ground thermal conductivity and borehole thermal  
396 resistance could be obtained from de data of table 3. The estimate of the ground thermal  
397 conductivity is the inverse of the slope so, considering the results of the three tests and  
398 averaging, the estimate will be  $\lambda=1,777\pm 0,007$  W/mK. And the estimate for the borehole  
399 thermal resistance is the intercept so, averaging,  $R_b=0,108\pm 0,012$  mK/W. It is remarkable that  
400 the three linear fits are very accurate with correlation coefficients very close to the unity. This  
401 fact is reflected in figure 11, in which the variables  $f_{0i}$  for the three tests calculated from the  
402 experimental data are shown (light grey line for the 1 kW test, dark grey line for the 2 kW test  
403 and black line for the 3 kW test), as the good agreement in the range of  $\tau-\Delta\tau$  starting in 0,03  
404 and finishing in 0,07 between the experimental points and the linear tendency described by the  
405 dashed grey line, plotted using the average of the slopes included in table 3, 0,563 mK/W.  
406 Nevertheless, this procedure is done for very early times and line source approach is valid for  
407 values of time much bigger than  $t_b = r_b^2/4\alpha$ , corresponding this value to  $t_b=0,74$  h, close to  
408 the fitting range, from 1,9 hours to 3,2 hours. So, this first estimation is not completely valid  
409 but serves to understand the further development.  
410



411 Fig. 11.- The values of  $f_{0i}=f_0-R_b0i$  against the values of  $\tau-\Delta\tau$  for the test  $i$  ( $i=1,2,3$ ) are represented.  
412

413 A qualitative analysis of the data presented in figure 11 drives to the following  
414 conclusions. First, after a  $\tau-\Delta\tau$  value of approximately 0,10 (corresponding to 4,64 hours) the  
415 data for the different tests starts to split, drawing curves in which the slope decrease with time  
416 (for long time periods) and this decrease depends on the thermal power injected, being this  
417 decrease greater when the injected thermal power is greater. This decrease can also be  
418 understood as dependency with the difference between the average temperature of the injected  
419 fluid and that of the undisturbed ground. Second, the three curves do not describe a linear  
420 behaviour in the variables represented as predicted by the line source approach. This fact can

421 be clearly seen comparing each experimental test curve with the dashed line representing the  
422 behaviour of the three tests at its beginning (values of  $\tau-\Delta\tau$  from 0,03 to 0,07). And third, local  
423 oscillations are observed in the three tests, short sections in which the slope can increase or  
424 decrease. The same oscillations are observed in the three tests, showing very similar patterns at  
425 the same values of  $\tau-\Delta\tau$ . Then, these oscillations are a physical effect, which can be interpreted  
426 as the appearance (conductivity increases, slope decreases) and disappearance (conductivity  
427 decreases, slope increases) of groundwater, phenomena activated by the injection of thermal  
428 power to the ground.

429 Although the qualitative analysis of the TRT data indicates slight discrepancies with  
430 line source prediction, it is interesting to analyse it with standard line source methodology. The  
431 data of the three tests have been fitted to the prediction of the infinite line source theory  
432 (equation 10) and to the prediction of the finite line source theory (equation 15) for different  
433 ranges of time. First range of time analysed is from  $t=10$  to  $t=40$  hours, starting from a value  
434 satisfying the constraint given in equation 3,  $t \gg t_b=0,7$  h, and for a time length of 30 hours.  
435 Estimates for the thermal conductivity and borehole thermal resistance for this fitting range,  
436 both analysis procedures and the three tests are included in table 4. First conclusion obtained  
437 after looking at these fitting results is that finite line source analysis estimates slightly lower  
438 values for both, thermal conductivity and borehole thermal resistance, than infinite line source  
439 analysis, with no significant differences. So, finite length effects could be neglected in this case.  
440 Second conclusion is that the three tests behave almost linearly in this fitting range, with  
441 correlation coefficients very close to the unity. Nevertheless, and as third conclusion, the  
442 estimates for thermal conductivity and borehole resistance differ considerably, being higher as  
443 the injected power increases. If the average of the three estimates is considered, and as error  
444 half the highest distance between them, then the prediction for thermal conductivity is  
445  $\lambda=2,445\pm 0,336$  W/mK (an error around 14%) and  $R_b=0,128\pm 0,011$  mK/W (an error around  
446 9%). Errors are associated to the discrepancy between the actual physical phenomena and the  
447 model that is not enough to describe them. This lack of accuracy reflects the fact that  
448 groundwater effects are already relevant in this fitting range, appearing in this line source  
449 methodology as an increasing dependency of the estimates for thermal conductivity and  
450 borehole resistance with the injected power. This dependency of the estimates does not mean  
451 that both depend on the injected power, it means that the line source approach is not able to  
452 describe groundwater effects and, then, these are artificially incorporated as estimates  
453 depending on input parameters as the injected power.

454 Second range of time analysed for tests 2 and 3 are from  $t=40$  hours until the end of the  
455 proof,  $t=98$  hours for test 2 and  $t=120$  hours for test 3. For test 1, a much longer experiment,  
456 two more ranges are analysed, a second range, from  $t=40$  hours until  $t=168$  hours, and a third  
457 range, from  $t=168$  hours until the end, in this case  $t=290$  hours. Estimates for the thermal  
458 conductivity and borehole thermal resistance for these fitting ranges, both analysis procedures  
459 and the three tests are included in table 4.

Fitting range	Test	Finite line source analysis			Infinite line source analysis		
		R <sub>b</sub> (mK/W)	λ(W/mK)	C.C.	R <sub>b</sub> (mK/W)	λ(W/mK)	C.C.
10-40 hours	1	0,107	2,065	0,972	0,108	2,093	0,972
	2	0,139	2,530	0,992	0,140	2,563	0,992
	3	0,139	2,738	0,981	0,140	2,774	0,981
40-168 hours	1	0,104	2,119	0,963	0,107	2,176	0,963
40-98 hours	2	0,147	2,768	0,955	0,149	2,831	0,954
40-120 hours	3	0,144	2,883	0,961	0,145	2,953	0,961
168-290 hours	1	0,099	2,117	0,855	0,104	2,208	0,854

461 Table 4.- Estimates for the thermal conductivity and borehole thermal resistance for several fitting ranges, finite line source  
 462 analysis and infinite line source analysis, and the three thermal response tests are included. Correlation coefficients (C.C.) of  
 463 each fitting are also included.

464 After looking at these fitting results for these second and third ranges, similar  
 465 conclusions as the ones reached for the first range are achieved. No significant difference is  
 466 found between estimates from infinite line source and from finite line source methodology.  
 467 Again, slightly lower estimates are obtained using finite line source analysis with very small  
 468 quantitative significance of the differences. The three tests still behave linearly but correlation  
 469 coefficients are lower than in first range, being more appreciable the discrepancies of the data  
 470 with the line source prediction. This means that groundwater effects also become higher with  
 471 time. Estimates for thermal conductivity and borehole thermal resistance achieve values higher  
 472 than in first range, so it seems that both parameters get higher with higher injected power and  
 473 with time.

474 From the qualitative and quantitative analysis of the three tests, two effects not described  
 475 by the line source approach are identified, both produced by groundwater advection  
 476 mechanisms. First one, observed for long time periods, is the increasing of the estimate for the  
 477 effective thermal conductivity as well as the borehole thermal resistance, being this increase  
 478 greater when the injected thermal power is greater. And second one, local oscillations in short  
 479 time ranges of the data, in which the slope, whose inverse is the estimate of thermal  
 480 conductivity, can increase or decrease. These short time oscillations can be interpreted as the  
 481 appearance (conductivity increases, slope decreases) and disappearance (conductivity  
 482 decreases, slope increases) of groundwater, phenomena activated by the injection of thermal  
 483 power to the ground.

484 One of the purposes of this research work is finding a phenomenological quantitative  
 485 description of the observed new effects. In particular, the long-term effect driving to the  
 486 unphysical result of a thermal conductivity depending on time and injected power. A  
 487 phenomenological parametrization of this effect can be obtained with an expression for the  
 488 effective thermal conductivity depending on the difference between the average fluid  
 489 temperature,  $T_{ave}$ , and the undisturbed ground temperature,  $T_0$ , as the following one:

$$490 \quad \lambda = \lambda_0 \left( 1 + x \frac{T_{ave} - T_0}{T_0} \right) \quad (16)$$

491 Where the new parameter,  $x$ , quantifies the effect produced by underground water  
 492 currents. Both observed dependencies, with time and with injected thermal power, can be  
 493 described with this approach. With this parametrization, groundwater effects are

494 phenomenologically integrated in the line source approach as a more complex definition of  
 495 effective thermal conductivity. Introducing this new definition for  $\lambda$  given in equation 16 in the  
 496 finite line source equation 15 it is obtained:

$$497 \quad [f_0 - R_b] \left(1 + x \frac{T_{ave}-T_0}{T_0}\right) = \frac{\tau-\Delta\tau}{\lambda_0} \quad (17)$$

498 Expression representing the new prediction of this improved line source approach for  
 499 the thermal response test data. Defining de quantity  $f_{GW}(x)$  as:

$$500 \quad f_{GW}(x) = [f_0 - R_b] \left(1 + x \frac{T_{ave}-T_0}{T_0}\right) \quad (18)$$

501 Final expression for this new prediction is formally identical to the previous one,  
 502 equation 15, just changing  $f_0-R_b$  by  $f_{GW}$ :

$$503 \quad f_{GW}(x) = \frac{\tau-\Delta\tau}{\lambda_0} \quad (19)$$

504 Note that this expression is developed on top of the finite line source approach but, if  
 505 infinite line source is the one involved, the same expression is valid just eliminating the term  
 506 including finite length effects,  $\Delta\tau$ .

507 Figure 12 is elaborated to check if this new expression predicts the experimental data  
 508 from the three thermal response tests. The quantity  $f_{GW}$  is plotted against the variable  $\tau-\Delta\tau$ , if  
 509 the prediction is correct then data will behave linearly and at the same position for the three sets  
 510 of data. To calculate the values for  $f_{GW}$  estimations for  $R_b$  and  $x$  are needed. In figure 12 values  
 511 for  $R_b$  are the same as the ones used in figure 11,  $R_{b01}=0,094$  mK/W,  $R_{b02}=0,117$  mK/W and  
 512  $R_{b03}=0,114$  mK/W, corresponding to a fit of the very early data of the test. Then, actual  
 513 definition for the vertical variable presented in figure 12 is:

$$514 \quad f_{GW_i}(x) = [f_0 - R_{bi}] \left(1 + x \frac{T_{ave}-T_0}{T_0}\right) \quad (20)$$

515 And the prediction to be observed:

$$516 \quad f_{GW_i}(x) = \frac{\tau-\Delta\tau}{\lambda_0} \quad (21)$$

517 Linear fittings of the variable  $f_{GW_i}(x)$  against the variable  $\tau-\Delta\tau$ , in the range between  
 518 10 hours and the end of each test (290 hours for test 1 kW, 98 hours for test 2 kW and 120 hours  
 519 for test 3 kW), for different values of  $x$  are performed. The one presented in figure 12  
 520 corresponds with the values of  $x$  minimizing the differences between the three estimates of  $\lambda_0$ .  
 521 The values obtained are  $\lambda_{0-1kW}=2,107$  W/mK,  $\lambda_{0-2kW}=2,166$  W/mK and  $\lambda_{0-3kW}=2,083$  W/mK for  
 522 the optimum value of  $x=3,4$ . Averaging the three values, an estimation for this parameter will  
 523 be  $\lambda_0=2,119$  W/mK, and taking as error half the maximum difference between them,  $\Delta\lambda_0=0,042$   
 524 W/mK. In long time periods, figure 12 show a clearer linear tendency of the variable  $f_{GW_i}(x)$   
 525 as a function of the variable  $\tau-\Delta\tau$ , in the time range analysed, 10 to 290 hours. This fact is  
 526 supported by the correlation coefficients of each fitting, very close to the unity, and better than  
 527 the ones obtained using line source approach (included in table 4). Values for correlation  
 528 coefficients are now: C.C.<sub>Test 1kW}=0,987, C.C.<sub>Test 2kW}=0,991 and C.C.<sub>Test 3kW}=0,991. Intercepts  
 529 of the fittings are slightly different from zero, then, a small correction to each value of the  
 530 estimate for borehole resistance is obtained, driving to the values  $R_{b1}=0,117$  mK/W,  $R_{b2}=0,141$   
 531 mK/W and  $R_{b3}=0,137$  mK/W. Averaging the three values an estimation for this parameter will  
 532 be  $R_b=0,133$  mK/W, and taking as error half the maximum difference between them,  $\Delta R_b=0,012$   
 533 mK/W.</sub></sub></sub>

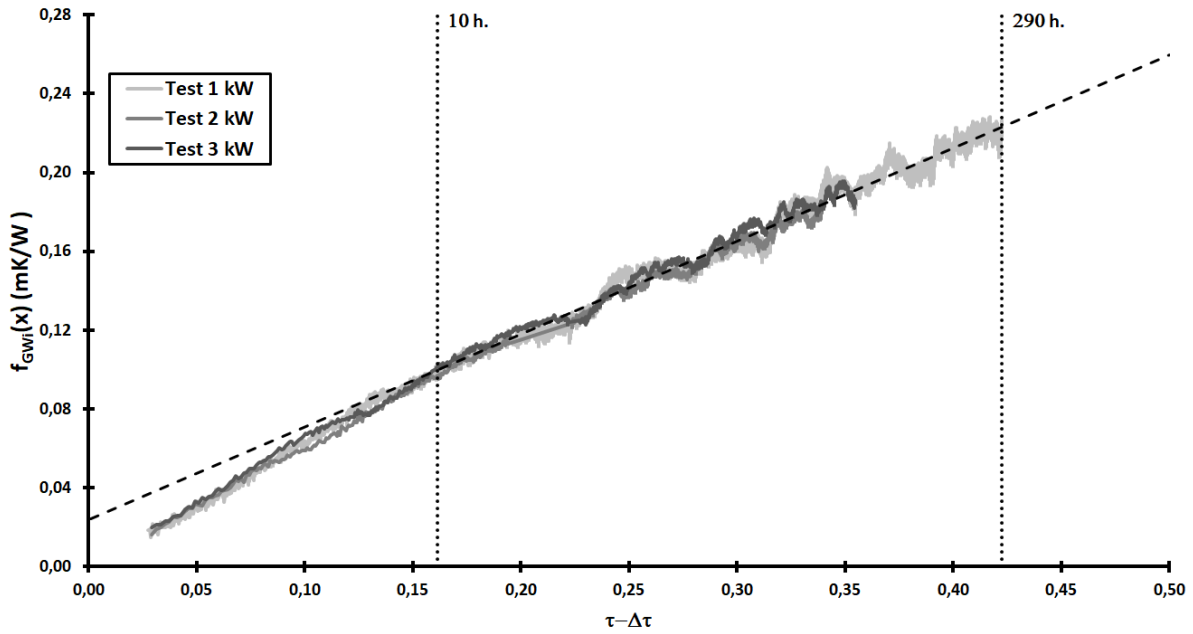


Fig. 12.- The values of  $f_{GWi}$  against the values of  $\tau-\Delta\tau$  for the test  $i$  ( $i=1,2,3$ ) are represented

534  
535

536 Together with the values of  $f_{GWi}(x)$  figure 12 also includes a dashed line representing  
537 the linear prediction obtained, line with slope  $m=1/\lambda_0=0,472$  mK/W and intercept  $a=0,024$   
538 mK/W. The agreement is quite well in the long range, describing appropriately experimental  
539 data, nevertheless, showing the short time oscillations with the really interesting fact that these  
540 oscillations are almost equal for the three tests when variables are presented as defined in  
541 horizontal and vertical axis of figure 12: experimental data of the three tests almost overlap.

542 The choice of the fitting range for this first analysis has been done with the criteria of  
543 starting at a point in which finite line source approach is valid and ending at the end of the  
544 acquisition period. Another interesting choice is a typical range of a standard test, starting at a  
545 point in which finite line source approach is valid and ending two days after. This fit has been  
546 also done, from 10 hours to 57,2 hours, including the results in table 5 (also in this table results  
547 for previous fit from 10 hours to the end are included). No significant differences between both  
548 fitting ranges are observed, being all estimates compatibles, and correlation coefficients very  
549 close to the unity.

550 Finally, another interesting choice for fitting range is an early starting (although is a  
551 range in which finite line source approach starts to be valid) because figure 11 showed that  
552 ground water effects appear even in this range. So, a fitting analysis starting in 1,9 hours has  
553 been done with to ends, one the end of the test and the other 57,2 hours. Results are included in  
554 table 5, showing, as expected, a slight decreasing of the estimate for parameter  $\lambda_0$ , but with very  
555 good agreement between the model and experimental results.

Finite line source analysis including ground water effects				
Fitting range	Test	$R_b$ (mK/W)	$\lambda_0$ (W/mK)	C.C.
10 hours-end	1	0,117	2,107	0,987
	2	0,141	2,166	0,991
	3	0,137	2,083	0,991
10 hours-57,2 hours	1	0,119	2,145	0,972
	2	0,139	2,117	0,991
	3	0,140	2,146	0,984
1,9 hours-end	1	0,114	2,066	0,991
	2	0,132	2,027	0,995
	3	0,131	2,001	0,995
1,9 hours- 57,2 hours	1	0,109	1,977	0,989
	2	0,128	1,954	0,996
	3	0,129	1,964	0,994

Table 5.- Estimates for  $\lambda_0$  and  $R_b$  using the finite line source analysis including ground water effects for the three thermal response tests and different ranges are included. Correlation coefficients (C.C.) of each fitting are also included.

556  
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559

## 560 5 DISCUSSION

561 The integration of the groundwater phenomenon in the finite line analysis model aims to  
562 reasonably describe the behaviour observed in the thermal response tests performed by the usual  
563 analysis methods. That is, a non-real increase in the value of the thermal conductivity depending  
564 on time and injected power, an augment that has also been assumed in the behaviour of the ground  
565 thermal heat capacity, thus keeping the thermal diffusivity constant. The introduction of a more  
566 complex definition of effective thermal conductivity in the model through a parameter that  
567 quantifies the effect produced by underground water currents allows to estimate the true value  
568 of the thermal conductivity,  $\lambda_0$ , regardless the power injected or the time elapsed in the test, as  
569 observed in the table 5. This is due to the fact that the effective thermal conductivity is divided  
570 into two terms, one static which is called “true value” because is unaffected by underground  
571 flow and another dynamic that depends on time and is characterized by parameter  $x$ . The  
572 differences in thermal conductivity values obtained both in different power injections and in  
573 different time intervals considered are less than 0,083 W/mK, that represents 4% of the average  
574 thermal conductivity value, while in the analysis performed through conventional finite line  
575 source method, table 4, this maximum difference is around 0,7 W/mK, that is 28% of the  
576 average thermal conductivity value estimated. This improvement in the results of the model  
577 analysis proposed is even more notable in the typical range time of a standard test, from 10  
578 hours to 57,2 hours, since in this case the difference between the values obtained represents  
579 1,4% of the average thermal conductivity whereas in conventional analysis, interval from 10 to  
580 40 hours, this difference is 27%.

581 This model has been applied using three injection pulses although, generally, thermal  
582 response test is performed only with an injection pulse but this does not exclude the application  
583 of this methodology because, in this case, the procedure to follow will be to find the parameter  
584  $x$  that maximizes the correlation coefficient. In this way, the objective of the work to simplify  
585 the analysis methodology for the thermal response tests performed under groundwater flow  
586 conditions intended for engineering applications is achieved.

587 In the analysis of the effective thermal conductivity value performed using traditional  
588 methodologies (finite line source and infinite line source analysis), figure 11 and table 4, it is  
589 noted that, in all cases, thermal conductivity value increases with the duration of the thermal  
590 power injection, being this augment smaller at low powers (test 1 kW). This phenomenon starts  
591 to notice after first hours of testing, becoming more important when 40 hours have elapsed.  
592 This is explained because the ground surrounding the borehole heat exchanger increases its  
593 temperature, as seen in the figures 5, 6 and 7, achieving a temperature gradient and activating  
594 the advective heat transfer mechanism. This is observed more clearly in layers composed by  
595 permeable materials such peat and gravel (around 10, 19 and 25 meters depth). In figures 8, 9  
596 y 10 an inverse phenomenon is detected during the ground recovery process, decreasing the  
597 temperature more quickly at depths mentioned above.

599 From the analysis carried out is deduced that, in the thermal response tests performed  
600 under standard conditions in locations with a groundwater presence, the effective thermal  
601 conductivity value is overestimated, which influences in the ground heat exchanger  
602 dimensioning. This positive effect will not always occur during the operation of the ground  
603 source heat pump installation because, among the factors that activate it, one of them is a  
604 continuous injection thermal power that not represents the usual heat pump operation,  
605 characterized by short work cycles. For this reason, the proposed analysis model is considered  
606 interesting because allows the designers choose the effective thermal conductivity value to  
607 which is more suitable to calculate with, the one that best represents the thermal conductivity

608 of the materials that make up the stratigraphic column of the borehole or the one that includes  
609 groundwater phenomena which helps heat transmission.

## 610 **6. CONCLUSIONS**

611 Detecting the effects of groundwater flow in the effective thermal conductivity is very difficult  
612 in standard thermal response test, obtaining a value which is not the thermal conductivity of the  
613 geological formation but a higher value. Existing models that consider convection required an  
614 extra resource effort with more complex data analysis, longer data collection periods or  
615 equipment to monitor more parameters.

616 In this work, a modification of finite and infinite line analysis models has been done in order to  
617 take into account the groundwater phenomenon in the thermal conductivity analysis. The  
618 implementation of this methodology has been possible thanks to an exhaustive thermal  
619 characterization of a borehole located at Universitat Politècnica of València, analysing the data  
620 of three different thermal powers injections during long periods of time in addition to a correct  
621 characterization of the ground undisturbed temperature and an analysis of the ground  
622 temperature behaviour before, during and after the injections. This analysis has allowed to  
623 identify the activation of the advective heat transfer mechanism explaining the increases  
624 detected in previous works in the value of the thermal conductivity depending on time and  
625 injected power. Using this background, a phenomenological parametrization of this  
626 groundwater effect in the thermal conductivity has been obtained introducing in the finite and  
627 infinite line source models an expression for the effective thermal conductivity that depends on  
628 a true thermal conductivity, the difference between the average fluid temperature and the  
629 undisturbed ground temperature and a parameter that quantifies the effect produced by  
630 underground water currents. The results show that it is possible to estimate the true thermal  
631 conductivity value regardless the power injected or the time elapsed in the test, fitting the values  
632 much better in the modified line source model than in the standard line source model.

633 So it can be concluded that, together with an adequate characterization of undisturbed ground  
634 temperature, which is already carried out in standard TRT, it is possible to analyse the effects  
635 of the groundwater flow without increase the TRT experimental measurements using the  
636 conventional analysis models, quantifying the “masking” of true thermal conductivity value  
637 due to groundwater flow action over time. Applying this methodology, GSHE designers can  
638 obtain the results of the geological effective thermal conductivity allowing a proper calculation  
639 without increase the cost of testing that benefits the development of GSHP installations.

640

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