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Additional Information

1 **Potential impacts of climate change on flow regime and fish habitat in mountain**
2 **rivers of the south-western Balkans**

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12

13

14 **Abstract**

15 The climate change in the Mediterranean area is expected to have significant impacts
16 on the aquatic ecosystems and particular in the mountain rivers and streams that often
17 host important species such as the *Salmo fariodes*. These impacts will most possibly
18 affect the habitat availability for various aquatic species resulting to an essential
19 alteration of the water requirements, either for dams or other water abstractions, in
20 order to maintain the essential levels of ecological flow for the rivers. The main scope
21 of this study was to assess potential climate change impacts on the hydrological
22 patterns and typical biota for a south-western Balkan mountain river, the Acheloos.
23 The altered flow regimes under different emission scenarios of the Intergovernmental
24 Panel on Climate Change (IPCC) were estimated using a hydrological model and
25 based on regional climate simulations over the study area. The Indicators of
26 Hydrologic Alteration (IHA) methodology was then used to assess the potential
27 streamflow alterations in the studied river due to predicted climate change conditions.
28 A fish habitat simulation method integrating univariate habitat suitability curves and
29 hydraulic modelling techniques were used to assess the impacts on the relationships
30 between the aquatic biota and hydrological status utilizing a sentinel species, the West
31 Balkan trout. The most prominent effects of the climate change scenarios depict
32 severe flow reductions that are likely to occur especially during the summer flows,
33 changing the duration and depressing the magnitude of the natural low flow
34 conditions. Weighted Usable Area-flow curves indicated the limitation of suitable
35 habitat for the native trout. Finally, this preliminary application highlighted the
36 potential of science-based hydrological and habitat simulation approaches that are
37 relevant to both biological quality elements (fish) and current EU Water policy to

38 serve as efficient tools for the estimation of possible climate change impacts on the
39 south-western Balkan river ecosystems.

40

41 **keywords**

42 Habitat series analysis, hydrological alteration, physical habitat simulation, West
43 Balkan trout.

44

45 **1. Introduction**

46 Several studies based on observations and modeling have pointed out that
47 hydrological systems and their biota are threatened from the indisputable fact of
48 climate change (Gedney *et al.*, 2006; Hauer *et al.*, 2013, Wu *et al.*, 2012; Zhou *et al.*,
49 2011). Still, the potential influence of recent climate fluctuations on the hydrological
50 balance of mountain riverine ecosystems has not been adequately studied. Mountain
51 rivers and streams are ecosystems with distinctive aquatic biota; these are of
52 outstanding value both for mountainous landscapes and for human activities and
53 economic development. Most mountain rivers are often located in poorly accessible
54 areas and are typically of small catchment dimensions, steep relief and high gradients
55 slopes. As a result of these characteristics, many mountain rivers have enjoyed the
56 environmental benefits of low or near-absent human impacts (Veza *et al.*, 2014),
57 although this is changing in the last few decades. Potential climate changes, altering
58 temperature and precipitation patterns may influence the hydrological balance of the
59 mountain riverine ecosystems leading to the limitation of available water resources for
60 all water users. In the Mediterranean basin for example, this water scarcity may be
61 especially acute during low-flow periods in summer. Vulnerable and habitat
62 specialized aquatic species such as salmonids may not be able to adapt to these
63 changes resulting in the degradation of ecological integrity of such rivers (Isaak *et al.*,
64 2010). Although several researchers have focused on the investigation of climate-
65 change effects on hydrological processes (Gibson *et al.*, 2005, Kim *et al.*, 2013; Luo
66 *et al.*, 2013), only recent studies have focused on researching the case of mountain
67 rivers (Beniston and Stoffel, 2014; López-Moreno *et al.*, 2014).

68 The mountain rivers of the south-western part of the Balkan Peninsula are
69 characterized by habitat heterogeneity which supports high species richness and
70 provide an area of international interest (Banarescu, 2004). Many of these upland
71 rivers maintain areas with natural and near-natural flow regimes and long-term

72 biogeographical isolation creates varied aquatic species assemblages (Skoulikidis *et*
73 *al.*, 2009; Zogaris *et al.*, 2009). Climate change impact studies for the mountain
74 waters of this region are poorly developed, although the importance of changes to
75 river flow regimes has been recently stressed (Angelini *et al.*, 2012; ENVSEC, 2012).

76 The main objective of this study was to assess potential climate change impacts on
77 the generic hydrological patterns and constituent fish habitats in a typical mountain
78 river system of the south-western Balkans focusing on a case-study in the upper part
79 of Acheloos River, Northwestern Greece. The actual near-natural status of the river
80 habitats for the West Balkan trout (*Salmo farioides*, Karaman 1938) in a
81 representative reach and the potential effects of climate change on the habitats of the
82 West Balkan trout (hereafter W. B.trout) were studied following the general principles
83 of the Ecological Limits of Hydrologic Alteration (ELOHA) framework (Poff *et al.*,
84 2010). In any similar study of potential changes in hydrology it is important to have
85 empirical evidence of the relationships between fish populations and their aquatic
86 habitats (Hauer *et al.*, 2013). W. B. trout was selected as a target species for several
87 reasons; it is an important indicator of high quality upland rivers, it dominates upland
88 cold-water streams (Economou *et al.*, 2007); it has a restricted distribution in upland
89 streams of the south-western Balkans, ranging from Montenegro to south-western
90 Greece (Kottelat and Freyhof, 2007); and it is assessed as a vulnerable species in a
91 state-wide species threat assessment (Zogaris & Economou, 2009). Furthermore,
92 salmonids play a crucial role in cold-water food webs and in the generation of
93 ecosystem services (Schindler *et al.*, 2010) and potential effects of climate change
94 both in terms of hydrological alteration and temperature may affect their habitats
95 (Almodóvar *et al.*, 2012).

96 To achieve the main objective, the following procedure was applied; i) Different
97 emission scenarios of the Intergovernmental Panel on Climate Change (IPCC)
98 obtained from regional climate models (RCMs) simulations were used to estimate
99 potential climate change impacts on flow regime using a hydrological model; ii)
100 following the streamflow alterations due to the changing climate conditions for the
101 different scenarios were assessed using the Range of Variability Approach (RVA), in
102 comparison with the simulated natural flow; iii) the physical habitat simulation
103 method integrating univariate habitat suitability curves and hydraulic modelling was
104 used to evaluate the plausible impacts on the relationships between hydrology and
105 biota using West Balkan trout as an indicator of biotic integrity.

106

107 **2. Materials and Methods**

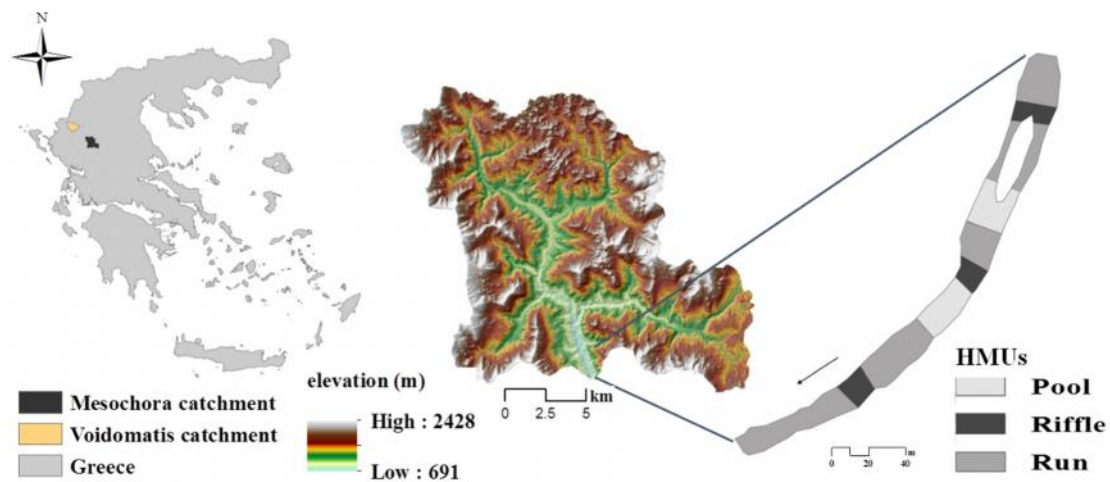
108 **2.1 Study area**

109 For this study, work was conducted in two river catchments in northwestern
110 Greece that show typical Mediterranean mountainous conditions widespread in the
111 south-western Balkans.; the Mesochora catchment in the upper Acheloos river, one of
112 the largest rivers in the Pindos Mountains; and the spring-fed section of the
113 Voidomatis tributary of the trans-boundary Aoos/Vjose river, near Greece's frontier
114 with Albania. The Acheloos' Mesochora catchment (632.8 km²) lies in the central
115 western mountainous region of Greece with mean elevation of 1390 m (Fig. 1). The
116 mean annual runoff of the catchment is 23.5 m³ s⁻¹ (Panagoulia, 1992). The mean
117 annual precipitation (weighted average over elevation bands) is 1898 mm. Most of the
118 precipitation falls between October and April (wet period) whereas at the higher
119 elevations the greatest amount of the precipitation falls as snow. The hydrology of the
120 Mesochora catchment is controlled by snowfall and snowmelt, with peak and low
121 flow occurring during May and September respectively. The water temperature in
122 summer ranges from 13.7 to 19 °C based on monthly measurements during June to
123 October 2013.

124 The particular catchments have been selected because they were relatively pristine
125 (close to reference conditions) since no significant water abstraction schemes and/or
126 pollution sources exist in the area. This was necessary in order to study the habitat
127 suitability and the impacts from hydrologic alterations for one of the most important
128 fish species (Western Balkan trout) of the area.

129 Habitat mapping of a 1.5 km river stretch of the upper Acheloos River (at 670 m
130 A.S.L, 39.479443°, 21.326510°, WGS 84) was carried out during low flow conditions
131 in the beginning of October 2013, in order to delineate the main features of the
132 physical habitat, based on field observations (Bisson *et al.*, 1982). More specifically,
133 identification of several types of HydroMorphological Units, hereafter HMUs (i.e.
134 pools, runs, riffles, glides, rapids), was made according to published methods (Dolloff
135 *et al.*, 1993), measuring their extent and physical attributes. Finally, a 390 m
136 representative river reach (Fig. 1), encompassing similar percentages and dimensions
137 of the surveyed HMUs, was selected as the representative reach (Mesochora reach).
138 The fish microhabitat-use survey, as part of the habitat simulation method was
139 conducted during summer 2014 in the Voidomatis River (39.948815°N, 20.693940°E,

140 WGS84). Voidomatis is a reference river with near-natural conditions within Greece's
141 Northern Pindos National Park. The catchment's mean annual precipitation typically
142 ranges between 1100 and 1700 mm, yielding a mean daily flow of $13 \text{ m}^3 \text{ s}^{-1}$
143 (Woodward *et al.*, 2008). Water temperatures in this karstic spring-fed stretch of the
144 river range from 10 to $12.5 \text{ }^\circ \text{C}$ based on field measurements during mid-summer
145 2014.



146

147 **Fig. 1.** Location of Mesochora and Voidomatis catchments (left), digital elevation
148 model of Mesochora catchment and distribution of hydromorphological units in the
149 representative reach of the Acheloos River (right).

150

151 2.2 Hydrological Model

152 In this study, the Soil and Water Assessment Tool (SWAT, Arnold *et al.*, 1998), a
153 process-based semi-distributed continuous hydrological model, was used for
154 simulating streamflow in the study area. SWAT has been successfully applied in
155 mountainous regions solving various environmental issues and exploring hydrological
156 fluxes (Abbaspour *et al.*, 2007; Debele *et al.*, 2010; Panagopoulos *et al.*, 2011).
157 Furthermore, SWAT has been used in many studies investigating climate and land use
158 change impacts on the water cycle and water quality (e.g. Ertürk *et al.*, 2014; Kim *et*
159 *al.*, 2013; Luo *et al.*, 2013; Salmoral *et al.*, 2015). SWAT is a process-based semi-
160 distributed continuous hydrological model. The watershed is subdivided into a set of
161 sub-watersheds connected with the river network. Each sub-watershed is further
162 divided into smaller basic units called Hydrological Response Units (HRUs), which
163 represent a combination of land use, soil and slope. SWAT simulates energy,
164 hydrology, soil temperature, mass transport and land management at HRU level.

165

166 2.2.1 *Model Setup, calibration and validation*

167 The main required spatial data for the parametrization of SWAT model is the
168 Digital Elevation Model (DEM), the land cover and the soil map of the catchment. In
169 this application, the topography was represented by a 25 m × 25 m DEM while
170 CORINE Land Cover (CLC), 1990 and 2000 databases, were used to represent land
171 cover. The soil information was derived by the European Soil Database (Panagos *et al.*,
172 2012) and by the geological maps of the National Institution of Geology and
173 Mineral Exploration (NIGME). Due to the data availability limitations both for the
174 model parametrization and the model calibration the results of SWAT application by
175 Panagopoulos *et al.*, (2011) in a nearby medium sized watershed were also taken into
176 account. The meteorological variables used to run the model were precipitation and
177 air temperature on daily time step. The required daily time-series of measured
178 precipitation and air temperature for three weather stations located inside the
179 watershed (“Katafyto”, “Pertoulio”, and “Theodoriana”) and one nearby station
180 (“Ioannina”) were provided by the Public Power Company of Greece (PPC) and the
181 Hellenic National Meteorological Service (HNMS) respectively. However, only
182 “Theodoriana” and “Ioannina” stations cover the entire simulated period (1983-2004)
183 and were used to run the model, while the other two stations were used to estimate the
184 precipitation and temperature lapse rates. The first three years of the simulation period
185 were used as a warm up period. Based on the DEM, the positions of the hydrometric
186 stations, and the location of the representative study reach (Mesochora reach), the
187 watershed was divided into 58 subbasins, and consequently into 2094 HRUs.

188 The calibration and validation were made at “Mesochora” gauging station for a
189 two-year period (October 1986 – September 1988) due to data availability limitations.
190 The first year was used for the calibration and the second year for the validation of the
191 model. Key considerations in the model calibration were the overall water balance and
192 the seasonal variation which were done at a monthly time step, as well as the low
193 flows which were done at a daily time step. Model performance was evaluated
194 statistically based on the Nash-Sutcliffe efficiency (NSE, – to +1, values close to +1
195 indicate better model performance) and the percent bias (PBIAS, indicator of under-
196 or over-estimation, values close to 0 indicate better model performance). The
197 performance was considered satisfactory if $NSE > 0.5$ and $PBIAS < \pm 25\%$ (Moriassi *et al.*,
198 2007; Rahman *et al.*, 2013). Specifically for the low flows, the model performance

199 was evaluated using the relative NSE (rNSE) and the NSE with logarithmic values
200 (lnNSE) to reduce the problem of the NSE sensitivity to extreme values (Krause et al.,
201 2005).

202

203

204 **2.3. Climate change scenarios**

205 The assessment of the potential regional changes in temperature and precipitation
206 patterns, under future emission scenarios (A1B for 2021-2050 and 2071-2100 time
207 periods and A2 for 2071-2100) of Intergovernmental Panel on Climate Change
208 (IPCC), was based on the study of Tolika *et al.*, (2012) who provided the
209 corresponding data for two annually distinguished periods.

210 Overall, the A2 and A1B are pessimistic scenarios in which the CO₂ concentration
211 will increase up to 815 ppm until the end of the century and on a global scale
212 temperature will rise from 2.5°C to 4.5°C. On the other hand the B2 scenario, a more
213 optimistic one, suggests a smaller mean planetary temperature rise ranging from 1.5
214 °C to 3.0 °C. All the scenarios are described in detail by Nakicenovic *et al.* (2000).

215 Tolika *et al.* (2012), considered twenty-two simulations from various Regional
216 Climate Model (RCMs) in order to assess the future changes in temperature and
217 precipitation with respect to the control period (1961-1990). All the models estimated
218 warmer and dryer conditions over the study area. For reasons of simplicity, in this
219 study the future changes in temperature and precipitation were considered by
220 adjusting the temperature and precipitation data series of the control period (1983-
221 2004) according to the average value of eight RCMs with resolution of 25 km under
222 A1B scenario for both 2021-2050 and 2071-2100 time periods and fourteen
223 simulations with a spatial grid resolution of 50 km for the period 2071-2100 under A2
224 (9 simulations) and B2 (5 simulations) scenarios (Table 1). A similar approach was
225 also used in previous studies (e.g., Jha *et al.*, 2006; Kalogeropoulos and Chalkias,
226 2012).

227

228

229

230

231

232

233 **Table 1**
 234 Rate of change in average precipitation (P) and temperature (T) projections based on
 235 RCMs (Tolika *et al.* 2012).

	Scenarios							
	A2		B2		A1B 2050		A1B 2100	
	P (mm)	T (o C)	P (mm)	T (o C)	P (mm)	T (o C)	P (mm)	T (o C)
Winter	-11.3	3.6	1.8	2.5	-3.7	1.4	-15.1	3.2
Summer	-53	4.8	-27.9	3.6	-19	1.9	-36.9	4.3

236

237

238 **2.4. Indicators of Hydrologic Alteration (IHA) and Range of Variability**
 239 **Approach (RVA)**

240 A common method to analyze the hydrological changes in a target river is to
 241 analyze streamflow time series with the indicators of hydrologic alteration (IHA),
 242 which allow the comparison between a baseline period (pre-impact) and another
 243 scenario (Richter *et al.*, 1996), in this case climate change scenarios (post-impact).
 244 The method relies on 33 parameters (median and coefficient of dispersion)
 245 corresponding to five fundamental characteristics of the flow regime (magnitude,
 246 frequency, duration, timing and rate of change) which greatly influence the ecological
 247 processes in river ecosystems (Poff *et al.*, 1997; Mathews & Richter, 2007).

248 The changes of the IHA were evaluated through the Range of Variability Approach
 249 (RVA) in the IHA software package (version 7.1; The Nature Conservancy, 2009). In
 250 a RVA analysis, the pre-impact data for each parameter (20 annual data or more) are
 251 divided into three categories of equal size; the *low* category with values lower or
 252 equal to the 33rd percentile; the *middle* category between the 34th and 67th percentiles;
 253 and the *high* category over the 67th percentile (default setup; The Nature Conservancy,
 254 2009). The program then compares the observed frequency of the values during the
 255 post-impact period with the expected frequency (pre-impact) of the IHA parameters
 256 within each of the three categories. The degree to which the RVA target category is
 257 not attained can be summarised in a hydrologic alteration factor (HAF), which is
 258 calculated for each of the IHA parameters as a percentage, that is:

259

260

261

262
$$\text{HAF}\% = \frac{O}{E} - \frac{f_i - E}{f_i} \quad (1)$$

263

264 Negative HAF values indicate that the frequency within a category will be
 265 decreased in the post-impact scenario. The value $\text{HAF} = -1$ corresponds to the
 266 condition when the event is not observed (in the specified category). A HAF is zero
 267 when the observed annual values under a scenario fall within the three RVA target
 268 ranges with the expected frequency (33% each category).

269

270 **2.5 Assessment of the available habitat for the West Balkan trout**

271 The flow requirements of three size classes of the W. B. trout have been assessed
 272 through the physical habitat simulation approach (Bovee *et al.*, 1998), in terms of
 273 depth and velocity, by combining Habitat Suitability Curves (HSC) and hydraulic
 274 simulation. During summer 2014 snorkelling was performed following international
 275 standards (Heggenes *et al.*, 1990; Martínez-Capel *et al.*, 2009) in the Voidomatis river
 276 to collect data on microhabitat-use by West Balkan trout; visual data were gathered
 277 for 103 large sized (>20 cm), 87 medium sized (10-20 cm) and 94 small sized
 278 (<10cm), individuals of W. B. trout. The HSC were developed following Bovee
 279 (1986); these curves relate the hydraulic or habitat variables with a suitability index
 280 (SI), ranging from 0 (unsuitable for the aquatic species) to 1 (excellent).

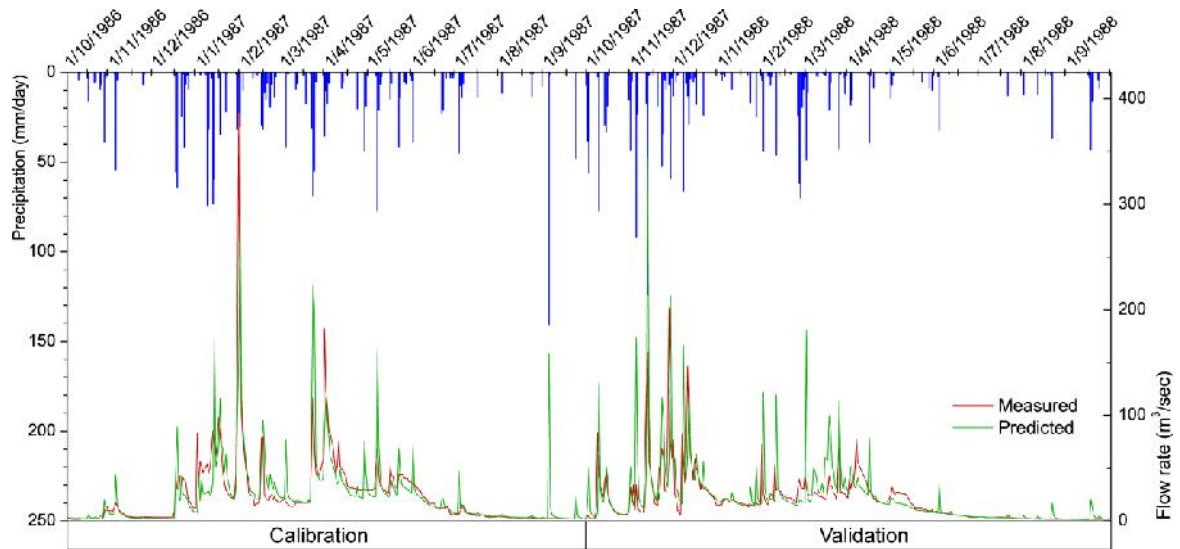
281 HEC-RAS (Version 4.1) was used to perform a pseudo-2D hydraulic simulation to
 282 estimate the changes in the depth and velocities for 30 stream flows, covering a wide
 283 range of possible summer flows under all the examined scenarios, ranging from 0.5 to
 284 $40 \text{ m}^3\text{s}^{-1}$ in the Mesochora reach. A topographic survey encompassing the main
 285 channel and banks was carried out with a GPS/GNSS Geomax - Zenith 20 using
 286 geodesic references (i.e. GGRS '87 – Greek Geodetic Reference System) to generate
 287 digital elevation models as the base for the model. Simulations were performed at 27
 288 cross-sections along the river reach. Every cross-section was subdivided in 12 cells
 289 both in the main channel and the overbank area and velocities calculated separately
 290 for each cell for the simulated water stage. In the physical habitat simulation, the
 291 hydraulic results were translated into the corresponding values of SI, through the use
 292 of the HSC, for each of the three size classes of fish independently. The geometric
 293 mean of the two SI for the hydraulic variables (i.e., the *combined* SI) was used to
 294 calculate the Weighted Usable Area (WUA) for every simulated flow. WUA is the

295 sum of the *combined* SI weighted by area, over all the points of the hydraulic model in
296 the Mesochora study site; this index was used as the general indicator of habitat
297 quality and quantity for each of the three sizes of W. B. trout, and the WUA-flow
298 curves were generated. In order to study only the suitable conditions for the target
299 species, WUA was estimated considering the cells with *combined* SI higher than 0.5
300 only (hereafter $WUA_{0.5}$). The whole procedure was carried out in R software (R
301 Development Core Team, 2012 R: A language and environment for statistical
302 computing). Moreover, habitat duration curves were constructed following the
303 procedures described within the IFIM methodology for environmental flow studies
304 (Bovee *et al.* 1998) indicating the exceedance probability for the potential habitat area in
305 the corresponding climate change scenarios with *combined* SI higher than 0.5.

306

307 **3. Results**

308 Due to data scarcity, the calibration in SWAT was based mainly on the curve
309 number (CN) parameter, while; specific attention was given to the calibration of the
310 base flow parameters alpha factor (ALPHA BF = 0.35) and lag (GW DELAY = 31)
311 due to their importance during habitat-limiting low flow periods and for
312 environmental flow regimes. The calibration of the base flow parameters was done
313 graphically at daily time step. The final step consisted of ensuring that the seasonal
314 flow balances were acceptable by slightly adjusting the estimated temperature lapse
315 rate (TLAPS) that significantly impacts the timing of snowmelt (TLAPS = 3.05
316 °C/km). Concerning the water balance and the seasonal variation (monthly time step),
317 the performance indicators in calibration were NSE = 0.69, PBIAS = 5.6%; and in
318 validation NSE = 0.51, PBIAS = 22.2%. Concerning the low flows prediction (daily
319 time step) the performance indicators in calibration were rNSE = 0.89, lnNSE = 0.85;
320 and in validation rNSE = 0.59, lnNSE = 0.96. The comparison between the simulated
321 and the observed hydrographs for both the calibration and validation periods is
322 illustrated in Fig. 2.



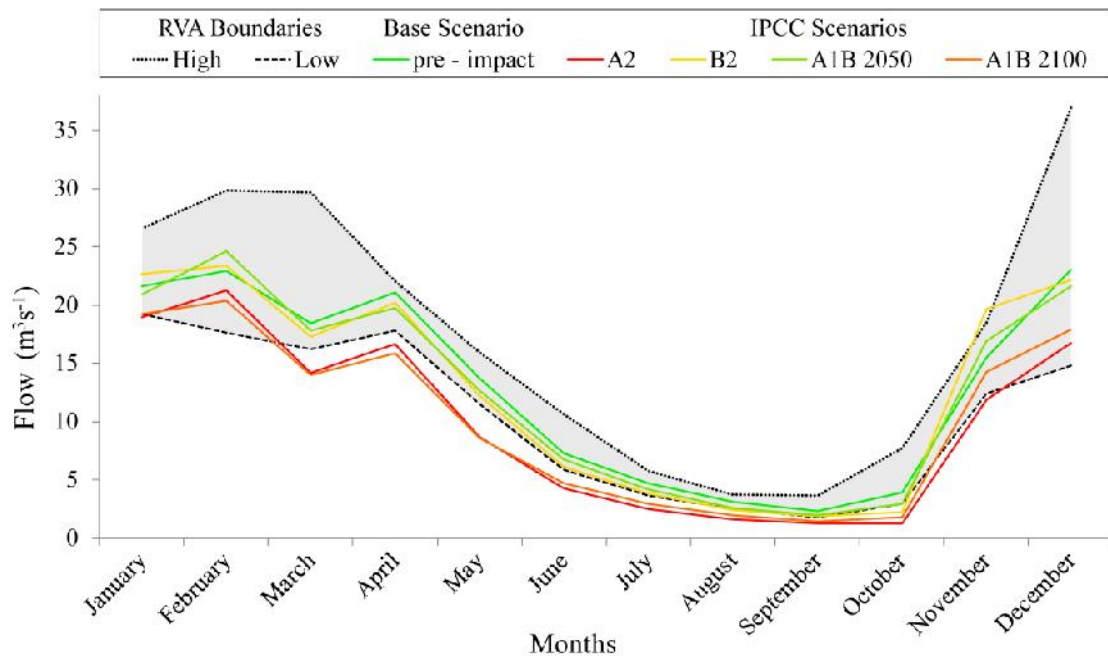
323

324 **Fig. 2.** Simulated and observed hydrographs for both the calibration and the validation
 325 periods.

326

327 The distribution of median monthly flows for the scenarios B2 and A1B-2050
 328 presented low hydrological alteration (Fig. 3). Conversely the A2 and A1B-2100
 329 produced lower stream flow in comparison with the middle category, especially
 330 during May (Table 2). However, to a certain degree, in the period from March to
 331 October all the projected scenarios presented lower monthly median flows than the
 332 pre-impact (baseline) period (Fig. 3). Specifically, the positive values of the
 333 Hydrologic Alteration Factors (HAF) in the *low* category (i.e. below the 33th
 334 percentile) for the months May to September corroborate this observation (Fig. 4).

335

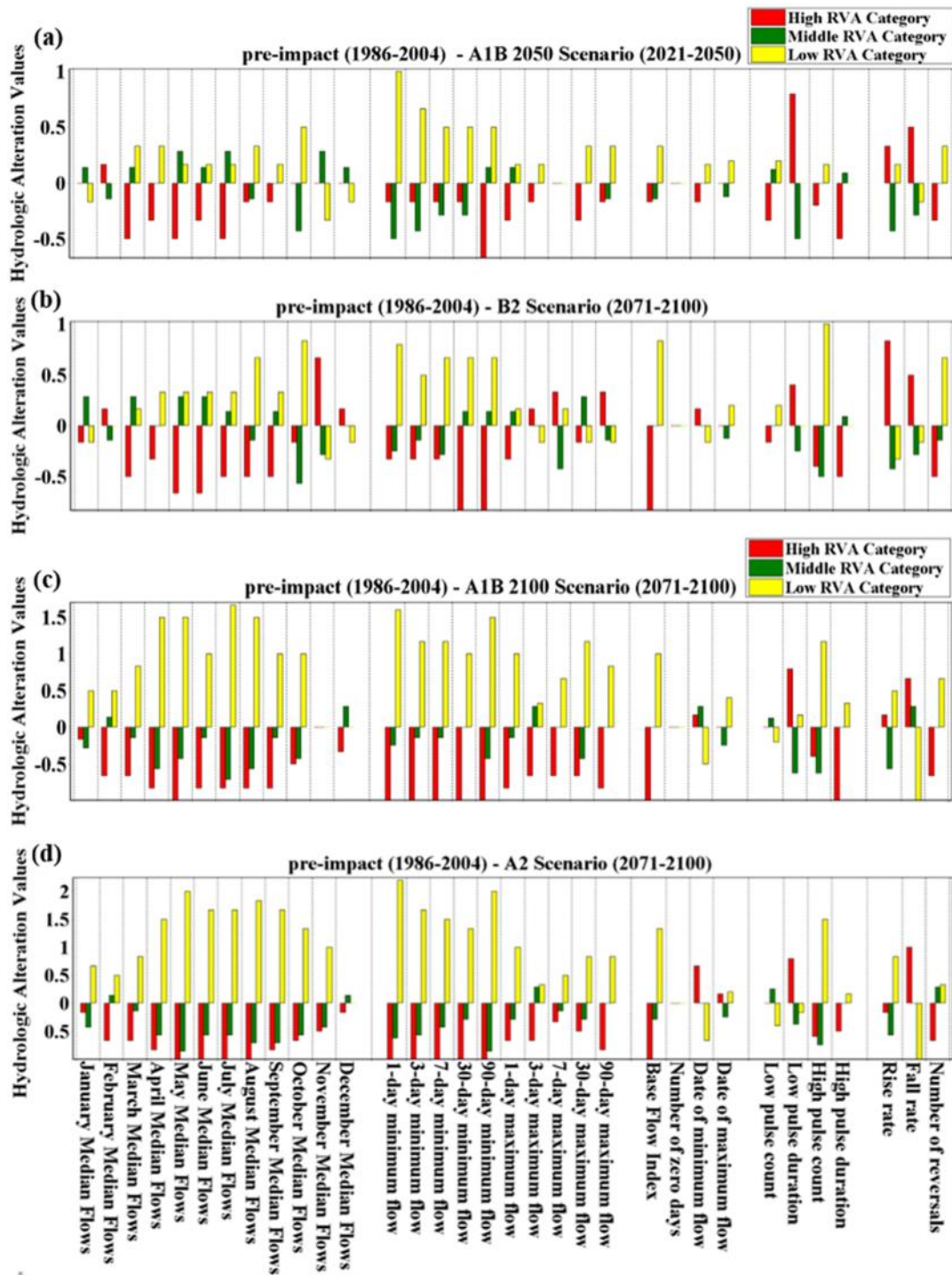


336

337 **Fig. 3.** Comparison of monthly median values of the projected scenarios with the
 338 middle category of the Range of Variability Approach – RVA (grey area within RVA
 339 boundaries).

340

341 The HAF related with minimum flows also showed a relevant increase in the low
 342 RVA category in all the scenarios, meaning that the minimum flows (several N-day
 343 minima) will be exacerbated (Fig.4); thereafter a high risk of droughts and limitation
 344 on water supply and suitable habitat is likely to occur. In addition, the low pulse
 345 duration was also increased (positive values in the high category) in all the scenarios.
 346 Regarding the maximum flows, in the A1B 2050 scenario (Fig. 4a) they will be
 347 slightly reduced (1-day, 3-day, 30-day, and 90-day maximum flow) but the scenario
 348 B2 lacks of a clear trend. The effects on high and low flows are more profound in the
 349 A1B 2100 and A2 Scenarios (Fig. 4c, 4d, respectively), as it is explained herein.



350

351 **Fig. 4.** Values of the Hydrologic Alteration Factors (HAF) comparing the pre-impact
 352 time period (1986–2004) and future climate change periods (a) A1B (2021 - 2050),
 353 (b) B2 (2071 - 2100), (c) A1B (2071 - 2100), (d) A2 (2071- 2100).

354

355

356

357 The scenarios A1B 2100 and A2 (period 2071 - 2100) presented high hydrologic
358 alteration. Regarding A1B 2100 (Fig. 4c) the positive values of the parameters related
359 to drought (1-day, 3-day, 7-day, 30-day, 90-day minimum) in the *low* category
360 indicated that droughts would occur more frequently. This fact is highlighted by the -1
361 values in the *high* category indicating the absence of the high flows events. Regarding
362 the parameters related to flood (1-day, 3-day, 7-day, 30-day, 90-day maximum)
363 negative values in the *middle* and *high* category indicate that in terms of magnitude,
364 the flood regime will be totally altered. On the other hand, the timing of the extreme
365 flow conditions, either low or high (group 3), will be only slightly affected.

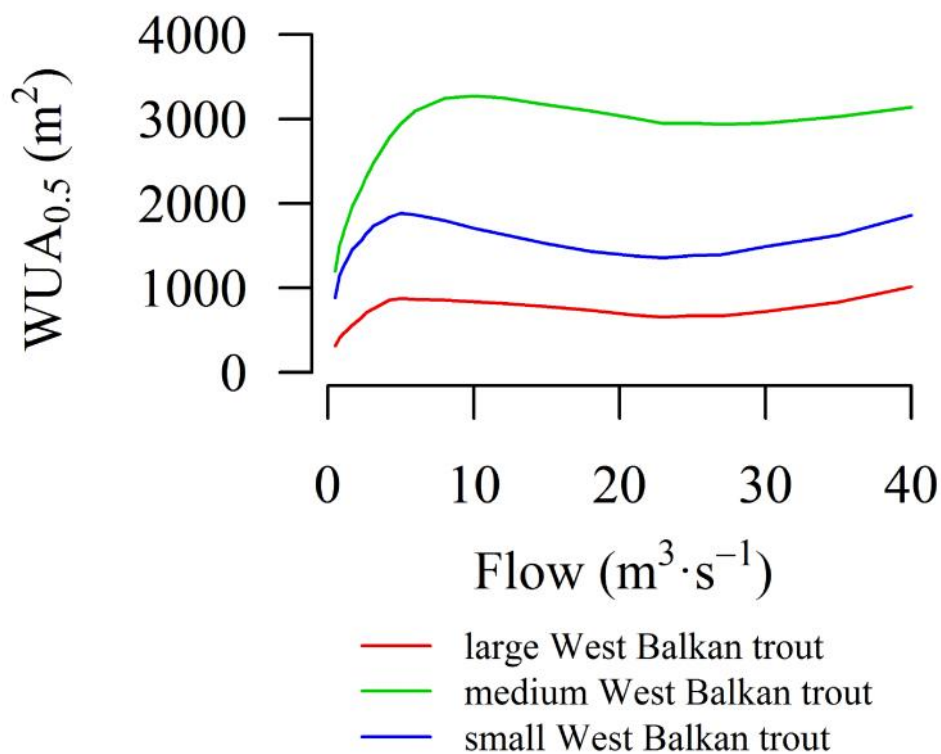
366 The A2 Scenario (Fig. 4d) was considered as the most altered because the *low*
367 category values were higher, for almost all the HAF, than in any other scenario in this
368 study. The HAF for monthly flows in the *middle* and *high* category were reduced from
369 April to November, expanding the dry conditions. This effect was also observed in the
370 reduction of HAF related to drought (1-day, 3-day, 7-day, 30-day, 90-day minimum)
371 where there is a complete absence of the *high* category values (-1), showing the
372 highest alteration from all the examined scenarios. A detailed comparison between the
373 simulated natural flow (pre-impact) and the worst case scenario (A2) with the median,
374 coefficient of dispersion and the HAF (version 7.1; The Nature Conservancy, 2009) is
375 presented in Table 2.

376

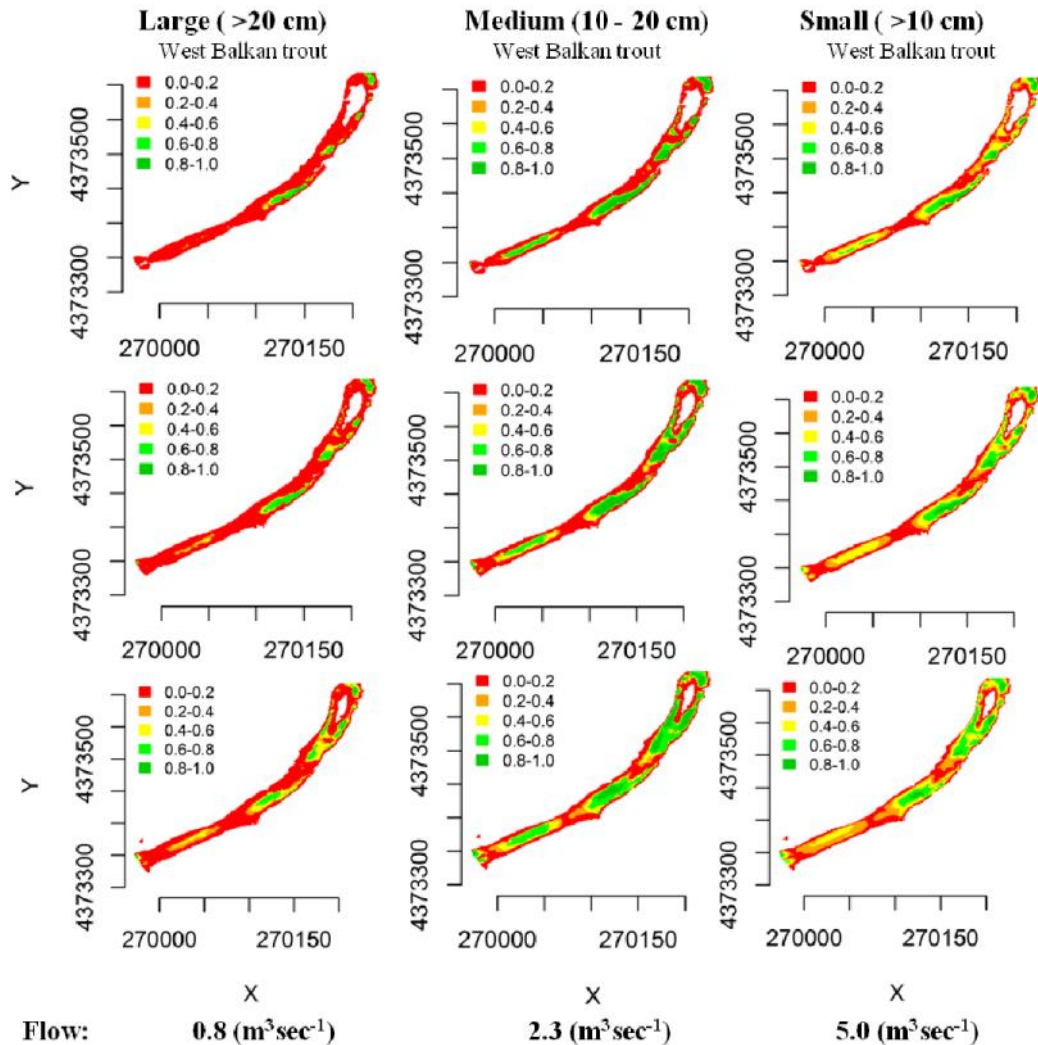
377 **Table 2**
 378 Hydrologic alteration analysis comparing Natural flow (pre-impact) with the A2
 379 (2071- 2100) Scenario.

	pre-impact		A2 Scenario		HAF		
	Medians	CD	Medians	CD	<i>High</i>	<i>Middle</i>	<i>Low</i>
Parameter Group 1 Flows (m ³ s ⁻¹)							
January	21.6	0.6	19.0	1.0	-0.2	-0.4	0.7
February	22.9	0.7	21.3	0.9	-0.5	0.0	0.5
March	18.4	1.1	14.2	0.9	-0.7	-0.1	0.8
April	21.1	0.4	16.6	0.5	-0.8	-0.6	1.5
May	13.7	0.5	8.7	0.5	-1.0	-0.9	2.0
June	7.3	0.9	4.3	0.5	-1.0	-0.6	1.7
July	4.7	0.5	2.5	0.5	-1.0	-0.6	1.7
August	3.1	0.7	1.7	0.5	-1.0	-0.7	1.8
September	2.3	1.2	1.3	0.7	-0.8	-0.7	1.7
October	3.9	1.9	1.3	1.8	-0.7	-0.6	1.3
November	15.6	0.8	11.9	0.8	-0.5	-0.4	1.0
December	23.0	1.3	16.8	1.6	-0.2	0.1	0.0
Parameter Group 2: Flow (m ³ s ⁻¹)							
1-day minimum	1.2	1.2	0.8	0.5	-1.0	-0.6	2.2
3-day minimum	1.3	1.3	0.8	0.5	-1.0	-0.6	1.7
7-day minimum	1.3	1.3	0.9	0.5	-1.0	-0.4	1.5
30-day minimum	2.1	1.2	1.0	0.8	-1.0	-0.3	1.3
90-day minimum	4.4	0.9	2.0	0.7	-1.0	-0.9	2.0
1-day maximum	547.7	0.7	413.8	0.6	-0.7	-0.3	1.0
3-day maximum	260.5	0.7	220.1	0.7	-0.7	0.3	0.3
7-day maximum	157.3	0.5	135.5	0.5	-0.3	-0.1	0.5
30-day maximum	77.8	0.5	64.1	0.4	-0.5	-0.3	0.8
90-day maximum	48.0	0.3	37.7	0.4	-0.8	-0.1	1.0
Number of zero days							
Base flow index	0.0	0.6	0.0	0.4	-1.0	-0.3	1.3
Parameter Group 3: timing of extreme water conditions							
Date of minimum	283.00	0.09	293.00	0.08	0.7	0.0	-0.7
Date of maximum	344.00	0.14	359.00	0.09	0.2	-0.3	0.2
Parameter Group 4: Frequency and duration of high /low pulses							
Low pulse count	7.0	0.6	9.0	0.3	0.0	0.3	-0.4
Low pulse duration	7.0	0.6	8.0	1.1	0.8	-0.4	-0.2
High pulse count	21.0	0.3	16.0	0.3	-0.6	-0.8	1.5
High pulse duration	2.0	0.5	2.0	0.5	-0.5	0.0	0.2
Parameter Group 5: Rate/ frequency of water conditions changes							
Rise rate	13.55	0.5	10.61	0.8	-0.2	-0.6	0.8
Fall rate	-0.61	-0.9	-0.34	-0.8	1.0	0.0	-1.0
Number of reversals	93	0.2	89	0.2	-0.7	0.1	0.5

381 The HSCs from the Voidomatis river indicated that large W. B. trout actively
 382 selected deep microhabitats (optimum; 1.4 – 1.8 m) with low velocities (optimum;
 383 0.15 – 0.30 ms⁻¹), whereas medium W. B. trout occupied medium-depth habitats
 384 (optimum; 0.60 – 0.95 m) and low velocity (optimum; 0.00 – 0.33 ms⁻¹). The small
 385 W. B. trout actively selected medium-depth habitats (optimum; 0.75 – 1.05 m) and
 386 low velocities (optimum; 0.00 – 0.30 ms⁻¹). The suitable areas summarized in the
 387 WUA_{0.5}-flow curves indicated low habitat availability for the large W. B. trout in
 388 comparison with the other two life stages; this observation applies to every simulated
 389 flow corresponding to the summer conditions under pre-impact and climate change
 390 scenarios, as shown in Fig. 5. To show the spatial arrangement of the suitable habitat,
 391 the maps for three representative summer flow values (0.8, 2.3 and 5 m³sec⁻¹) are
 392 depicted in Fig. 6.



393
 394 **Fig. 5.** Curves relating WUA_{0.5} and stream flow within a range of summer flows
 395 which are possible to occur under all the examined scenarios for three size classes of
 396 W. B. trout.
 397



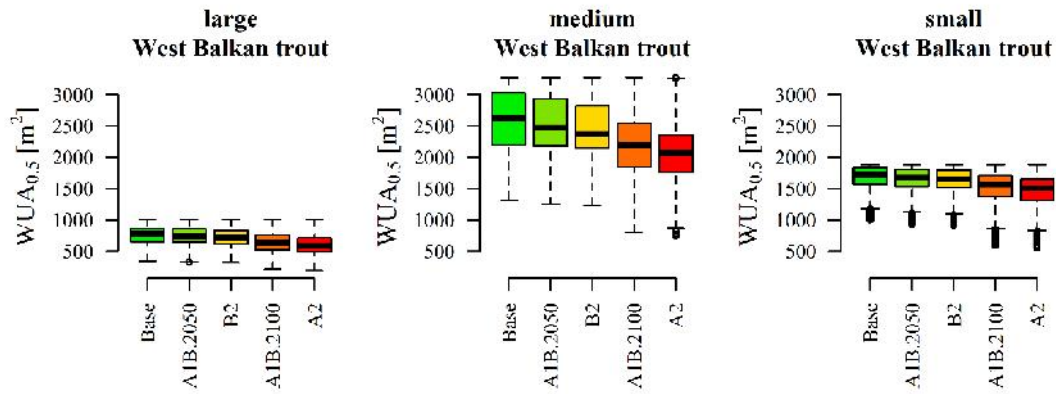
398

399 **Fig. 6.** Habitat suitability maps depicting the *combined* SI of the study area for three
 400 flows and the three size classes of the West Balkan trout.

401

402 Based on applied scenarios, the study streams are sensitive to climate fluctuations.
 403 The habitat analysis presented a similar pattern to the hydrological analysis *via* the
 404 RVA method. Therefore, the scenario corresponding to the lowest alteration was the
 405 A1B 2050, whereas the worst scenario was A2. The most affected size class would be
 406 the medium W. B. trout, both regarding the magnitude and frequency followed by the
 407 small W. B. trout (Fig. 7). Finally the size class which will be less affected would be
 408 the large W. B. trout since it presented a small suitable area for any analyzed flow
 409 (Fig. 7).

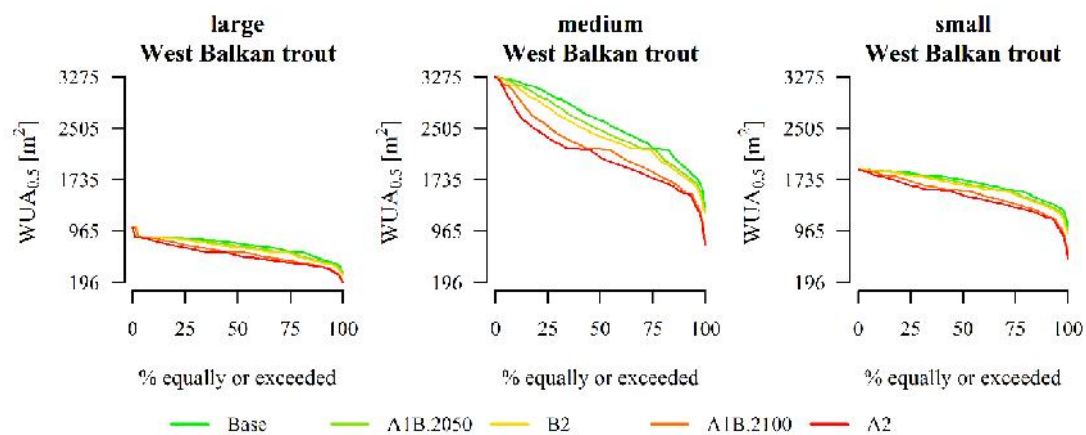
410



411
 412 **Fig. 7.** Box-plots showing the projected changes in $WUA_{0.5}$ for three size classes of
 413 the West Balkan trout under summer conditions (July to August) for the pre-impact
 414 period and the examined climate changed scenarios.

415
 416 The differences in habitat projected for the summer period were mainly caused by
 417 reductions in water depth. These reductions were especially important under the A2
 418 and A1B 2100 scenarios (Fig. 8). Consequently, the habitat duration curves are much
 419 lower than in the other two scenarios, thus suggesting a significant degradation of the
 420 suitable habitat area for the W. B. trout. However none of them were low enough to
 421 suggest the extirpation of the species of the study site because the minimum $WUA_{0.5}$
 422 was in any case larger than zero. The general reduction in water resources is translated
 423 to river habitats where the reduction was observed in events of any frequency and
 424 magnitude.

425



426
 427 **Fig. 8.** Habitat duration curves (HDC) for the comparison between the pre-impact
 428 conditions (baseline) and the expected ones under the four examined scenarios in terms of
 429 probability of exceedance.

430 **4. Discussion**

431 **4.1 Hydrologic modelling**

432 Streamflow simulation is often challenging in mountain river catchments because of
433 high relief topography and complex hydrological processes. Rates of change in
434 precipitation and temperature with respect to elevation and strong spatial variability of
435 meteorological conditions often limit the ability to accurately reproduce stream runoff
436 by hydrological models (Rahman *et al.*, 2013; Soulis and Dercas, 2007). Furthermore,
437 in many cases, especially in the less developed regions, the meteorological
438 information available is scarce and confined to lower altitudes or coastal locations
439 (Brito *et al.*, 1999; Soulis, 2015). This problem is further exacerbated by limited site-
440 based hydrological and environmental data availability, which is especially the case in
441 the Balkan countries (Skoulikidis *et al.*, 2009). Nevertheless, despite the data
442 limitations, the model performance was considered acceptable according to the
443 criteria posed (Moriassi *et al.*, 2007; Rahman *et al.*, 2013), and presented similar
444 values in comparison with previous studies performed in Greek river basins
445 (Gamvroudis *et al.*, 2015) thus highlighting the validity of any further analysis.
446 Especially for the case of low flows, which are of particular importance for the scope
447 of this study, the model performance was much better and it was considered
448 satisfactory as well. Therefore, even if the remaining uncertainty is an important
449 constrain, the overall model performance was considered adequate for the purposes of
450 a comparative analysis given the data scarcity that characterizes the study area.

451

452 **4.2 Potential climate change impact on freshwater mountain river systems**

453 Results gained in this study show that the A1B 2050 and B2 scenarios have limited
454 impact in comparison with the other two scenarios (A1B 2100 and A2) where
455 reductions in the precipitation during winter period and temperature increments
456 during summer period affect streamflow, especially by reducing the magnitude and
457 increasing duration of low flows. Our results corroborate a broad scale analysis on the
458 expected impacts of the different climate change scenarios that already suggest the
459 major impact of the A2 scenario in river flows (Van Vliet *et al.*, 2013). Moreover, in a
460 similar study in Spain, Salmoral *et al.* (2015) concluded that increasing mean
461 temperature is the main factor supporting increasing evapotranspiration and thus
462 driving streamflow reduction. Studies in Mediterranean-climate streams observed a
463 lack of resilience and negative impacts to biodiversity due to prolonged droughts

464 related with long-term habitat changes induced by the increment in frequency and
465 magnitude of the low flow events (Bêche et al., 2009). Moreover innovative
466 approaches, such as microsatellite DNA analyses, revealed that the effects of drought
467 may be profound and long-lasting, resulting in population bottlenecks and altering the
468 course of the evolution of species (Humphries and Baldwin, 2003).

469 Mountain streams in Mediterranean regions have highly variable seasonal
470 discharge patterns, with torrential flood pulses and seasonal drought periods, being
471 usually much less flashy and more variable than temperate stream systems (Bonada *et*
472 *al.*, 2007). The areas chosen for this study represent some of the most natural
473 mountain river corridors as has been shown both by instream studies (Chatzinikolaou
474 *et al.*, 2006; Economou *et al.*, 2007) and riparian corridor assessments (Zogaris *et al.*,
475 2008); the study areas therefore provide excellent baselines to study biotic-abiotic
476 interactions in near natural states before potential climate-driven changes take place.
477 The south-western Balkan region, encompassing the Adriatic and Ionian basins, has a
478 humid Mediterranean climate and receives a much higher precipitation compared to
479 the eastern Balkans; it has distinctive aquatic biocommunities with a very high
480 proportion of endemic species, being a biogeographically isolated region and a
481 refugium area during the Pleistocene glaciations (Zogaris *et al.*, 2008; Skoulikidis *et*
482 *al.*, 2009).

483 The study focused in the W. B. trout although the whole ecosystem is likely to be
484 affected (Bêche *et al.*, 2009; Humphries and Baldwin, 2003; Mantua *et al.*, 2010;
485 Wenger *et al.*, 2011). The habitat alterations predicted in this study have been known
486 to cause geomorphic simplification, floodplain disconnection and disruption of lateral
487 and longitudinal connectivity, thereby affecting habitat dynamics and making it
488 difficult for native biota to adapt (Poff *et al.*, 2007). The prescribed scenario changes
489 will not only affect instream biota but most probably the area's riparian vegetation,
490 which currently supports species-rich near-natural floral assemblages (Zogaris *et al.*,
491 2008). However, in many cases of mountain rivers there is a significant scarcity of
492 historical flow data and very limited information about the flow requirements of the
493 river's biota. Furthermore, the overfishing that is taking place especially in the recent
494 years may influence predictions, specifically in the larger size class trout.
495 Consequently, climate change together with inadequate water management along with
496 the insufficient water conservation policy are now interpreted as major threats for
497 mountain streams, leading to alterations which may rapidly degrade ecosystem

498 structure and ecological processes and the services they provide (Postel & Richter,
499 2003). Moreover, the negative impacts of climate change are projected to be most
500 pronounced particularly in relatively pristine, high-elevation and headwater streams
501 where restoration measures are usually not possible (Battin *et al.*, 2007). This
502 situation is especially sensitive in the mountain rivers of the south-western Balkans
503 where there are severe problems with changes in land-use and poor management of
504 water resources, dam developments, pollution control and protected area development
505 (Chatzinikolaou *et al.*, 2006; ENVSEC, 2012).

506

507 **4.3 Relationships of climate-impacted flow regimes and fish habitat**

508 Using the native trout species as a habitat-specialized indicator is a practicable
509 application because this species is phylogenetically related to the well-studied cold-
510 water specialist, the European brown trout (*Salmo trutta* L.) often dominating
511 mountain stream waters (Economou *et al.*, 2007; Kottelat and Freyhof, 2007). Broad-
512 scale studies of climate change effects on freshwater species have traditionally
513 focused mainly on temperature, underrating critical drivers such as flow regime and
514 biotic interactions (Wenger *et al.*, 2011).

515 In the Mediterranean context previous efforts to quantify the expected effects of
516 climate change on cold water salmonids (European brown trout) stated temperature
517 alteration as the main driver for the expected shrinkage of the trout distribution area
518 (Almodóvar *et al.*, 2012) whereas other studies, in colder climatic conditions,
519 considered the flow as the keystone to assess the impact of climate change on the
520 distribution area of salmonids (*Oncorhynchus spp.*) (Wenger *et al.*, 2011). Our results
521 were in line with the latter study, suggesting that the alteration of stream flows,
522 especially by reducing them, will turn in a reduction in the suitable habitat available
523 for the target species. Furthermore, as it has been pointed out in other studies (Mantua
524 *et al.*, 2010), it is very likely that more stable flow patterns and reduced flow predicted
525 by our analysis may increase temperatures in our upland Mediterranean mountain
526 streams, thus producing additive pressure against the survival of cold-water species
527 such as the native salmonid. Regarding physico-chemical conditions of the aquatic
528 habitat, it is also relevant to consider that a reduction of water quality, which would
529 produce severe risks for the ecosystem integrity, is probable under future scenarios of
530 water scarcity in Mediterranean rivers (Petrovic *et al.*, 2011).

531 The specific results of the habitat analysis showed that the habitat suitability for W.
532 B. trout will suffer the effects of hydrological changes with the A2 scenario
533 corresponding to the worst situation for this keystone fish species. Furthermore, as it
534 is also pointed out by Hauer *et al.* (2013) by analyzing the impact of altered stream
535 flows in a smaller scale, the increased frequency of low flows, especially in the
536 summer periods, will reduce habitat quality and quantity as a result of the changes in
537 depths and velocities. The Mesochora reach resembles a relatively deep run with high
538 velocities and medium depths; for this reason the potential impacts of climate change
539 seem to induce relatively small changes on the quality of the habitats for small and
540 large size trout, since the physical characteristics were out of their preference range
541 (i.e. suboptimal habitat conditions for this species during summer flows). However, as
542 well as the reduction of water resources at a monthly time scale, the reduction in the
543 magnitude and frequency of high-flow events was consistent among scenarios, with
544 severe potential impacts during the winter, when these events should prepare
545 spawning habitats and provide with cues for the salmonid populations before the
546 migration and spawning. Therefore, according with previous studies, we can
547 hypothesize a very relevant impact of the smaller frequency of peak flows, producing
548 a reduction in fish recruitment, turning into decreased abundance or extirpation of
549 native fishes in the long term (Mathews & Richter, 2007; Poff *et al.*, 2010).

550 Thereby small trout selected relatively shallow and slow flow microhabitats, which
551 were scarce in the study site, whereas large trout selected preferably deep and slow
552 flow microhabitats (i.e. pools) which were, likewise, absent in the study site. As a
553 consequence of the velocity-depth distribution the Mesochora study site became
554 especially suited for medium size W. B. trout and then the forecasted reduction in the
555 running flows will mostly affect this size class. Although in a different magnitude, the
556 negative impact of the flow reduction affected all size classes. We have not predicted
557 null WUA_{0.5} for any considered scenario and class but the long term effect is likely to
558 reduce the presence of the species in the study area which is already affected by the
559 severe overfishing, even involving illegal spear fishing and electrofishing imperiling
560 the W. B. trout populations. Furthermore we considered the W. B. trout the target
561 species neglecting the effect on the rest of the ecosystem when it is well documented
562 that flow regime and temperature but also biotic interactions can drive differential
563 declines of trout species under climate change (Wenger *et al.*, 2011) thus composing a
564 synergic suite of factors that could lead to species extirpations (Brook *et al.*, 2008).

565

566 **4.4 Conclusions**

567 This study suggests that a changing climate may affect instream flow conditions in
568 mountain rivers which will probably impact ecological integrity. In the Upper
569 Acheloos river, as in many other mountain rivers in the southwest Balkans, there are
570 many influences affecting the flow regime as well as other essential habitat
571 characteristics (Mathews and Richter, 2007; Skoulikidis *et al.*, 2009). Effective
572 conservation of biodiversity in mountain river systems requires accurate downscaling
573 of climatic trends to local habitat conditions. Nevertheless, downscaling is difficult in
574 complex and diverse terrains, as those of the mountains, with varied microclimates
575 and special local characteristics (Isaak *et al.*, 2010). Our study is one of the few
576 attempts to use a prominent indicator fish species to explore specific habitat changes
577 based on its studied requirements within reference mountain river stretches.

578 The IHA analysis of the Upper Acheloos River indicates that the flood regime
579 under all the examined scenarios, including both small and large floods, will be
580 altered. The physical habitat simulation method suggested severe changes on the
581 habitat quality and quantity since impacted flow regimes showed a lack of flood-pulse
582 peaking and reduced water quantity. This important element may lead to warmer more
583 stable conditions and some microhabitats required by rheophilic cold-water fauna may
584 show marked decline. According to our results from the current case-study, the most
585 affected trout size class will be the middle sized, then the small and finally the large.
586 However apart from the empirical evidence, it is probable that the large-sized trout in
587 the Acheloos have suffered severe declines from overfishing, and in this case there is
588 still uncertainty for this size-class category. Flow-ecology relationships need further
589 investigation by identifying specific hydrologic alterations that may impact particular
590 species, biocommunities and ecological processes.

591 To refine predictions based on climate change instream alterations, further research
592 should be conducted to understand the mechanisms associated with the biological
593 responses to the climate effects. Possible interactions between climate change, water
594 quality, and food availability due to ecosystem changes; fragmentation of species
595 populations due to thermal constraints; increases in predation; and changes in species
596 interactions and competition within aquatic ecosystems should be analyzed towards a
597 holistic approach. Finally, future changes in other anthropogenic stresses on fish
598 habitat, such as increasing water withdrawals, dams or changing land use must also be

599 quantified and analyzed. Thereafter, more research is needed to investigate the effects
600 of flow reduction and flow regime change on the instream environments of mountain
601 rivers (Dewson, 2007). This is especially important in sensitive areas with high-
602 endemicity aquatic ecosystem within restricted freshwater ecoregions such as in the
603 southwestern Balkans.

604

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